

# Overview of Data Assimilation Methods

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# Outline

- What is data assimilation and how does it work?
- Observations and model error sources
- Blending observations and model: the Bayes perspective
- Particle Filters
- Kalman Filter methods
- Variational methods
- Hybrid data assimilation

# Data Assimilation

NWP definition: Process by which “optimal” initial conditions for numerical forecasts are estimated.

- The best analysis (initial conditions) is the analysis that leads to the best forecast
- Optimal in a statistical sense: minimises error and/or maximises probability of the analysis
- Do it quickly – typically in less than 45 minutes on a large high performance computer (for global NWP; for local area NWP this can be much less!)

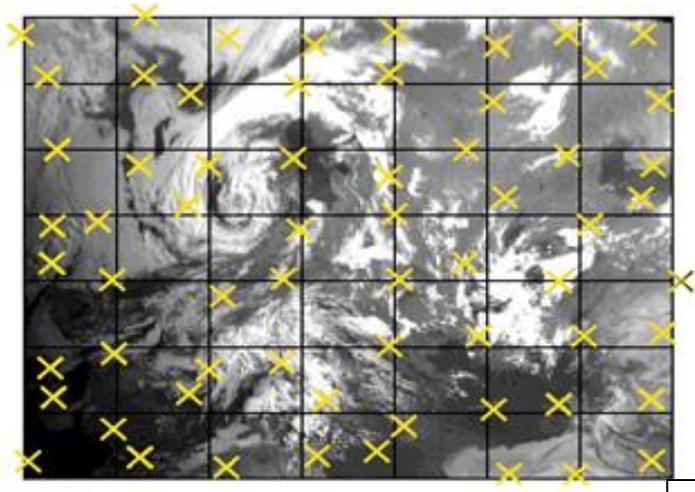
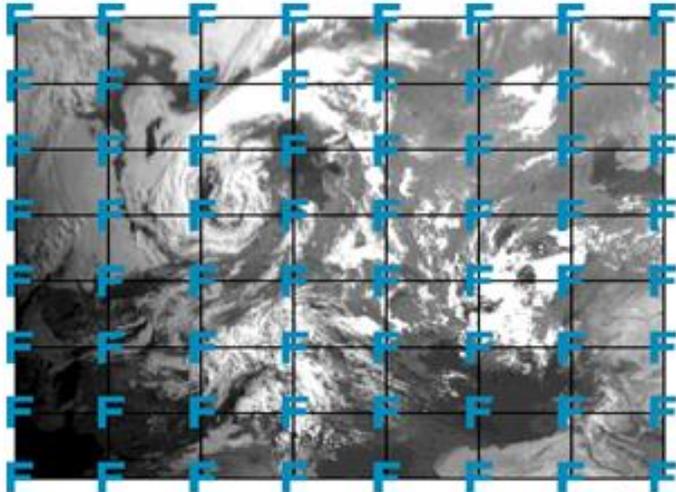
# Data Assimilation

Data Assimilation has two main goals:

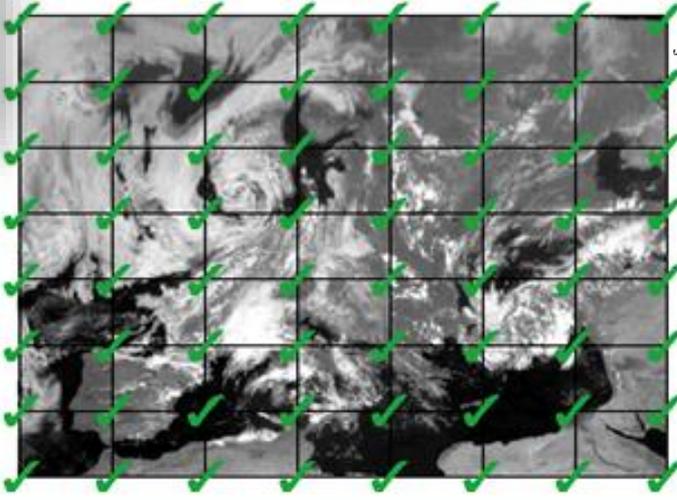
- Provide the **best estimate** of the initial state of the atmosphere-land-ocean system out of all available information (model + observations)
- Quantify the **uncertainty** of our estimate of the initial state (this is necessary to be able to initialise an ensemble forecast!)

Model Forecast (with errors)

Observations (with errors)



Computer (with a lot of CPUs)

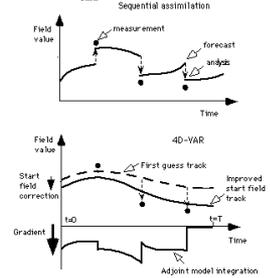


Analysis (with - smaller - errors)

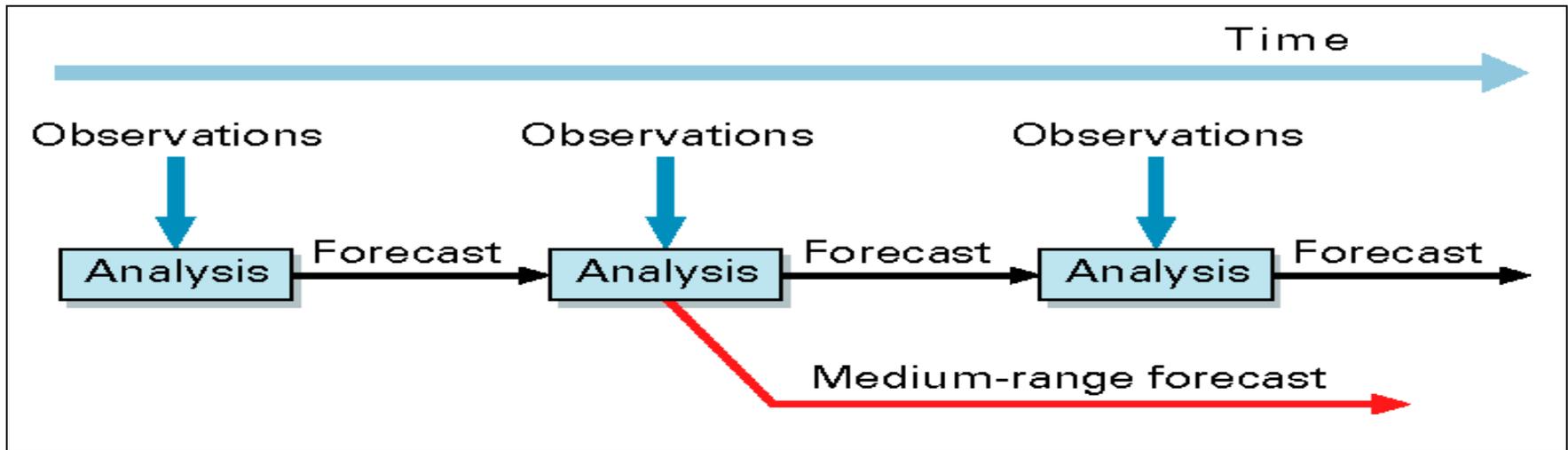


People (with a lot of good ideas)

$$J(x) = (x - x_b)^T B^{-1} (x - x_b) + \sum_{i=0}^n (y_i - H_i[x_i])^T R_i^{-1} (y_i - H_i[x_i])$$

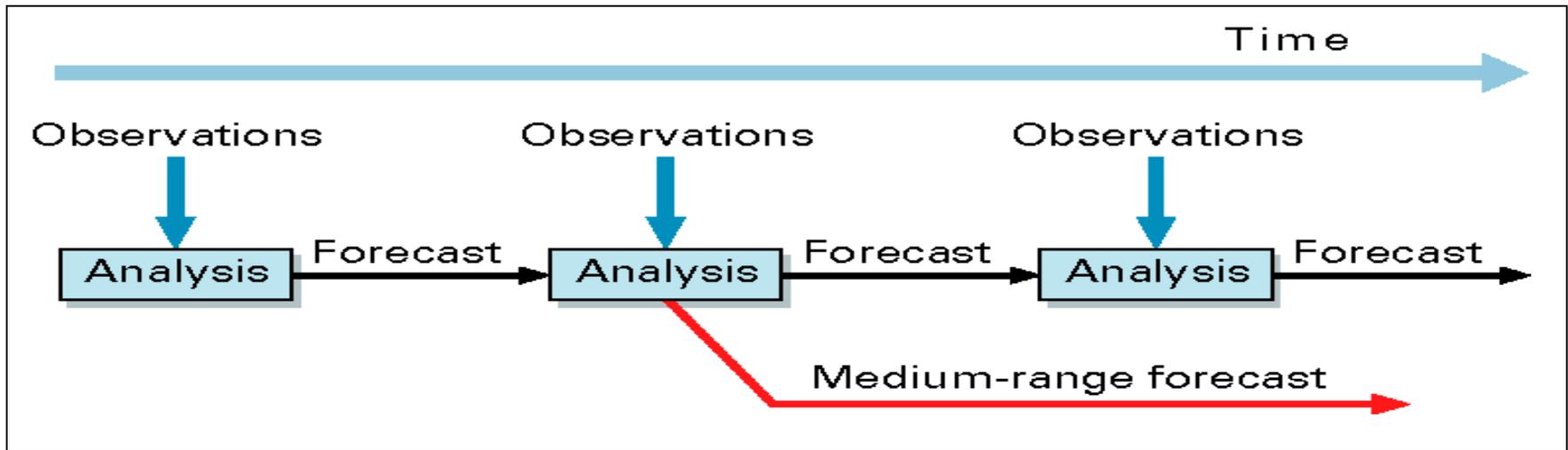


# The Data assimilation cycle



- An analysis is not produced by observations alone!
- The observations are used to correct errors in the short forecast from the previous analysis time (every 12 hours at ECMWF, 6-hourly in other global NWP; more frequently for higher resolution, local area models).
- The short range forecast carries information from past observations into the current analysis

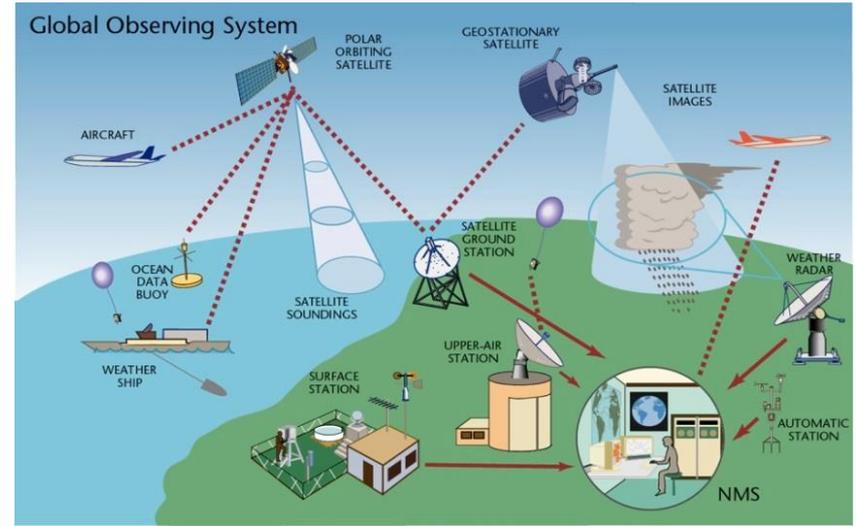
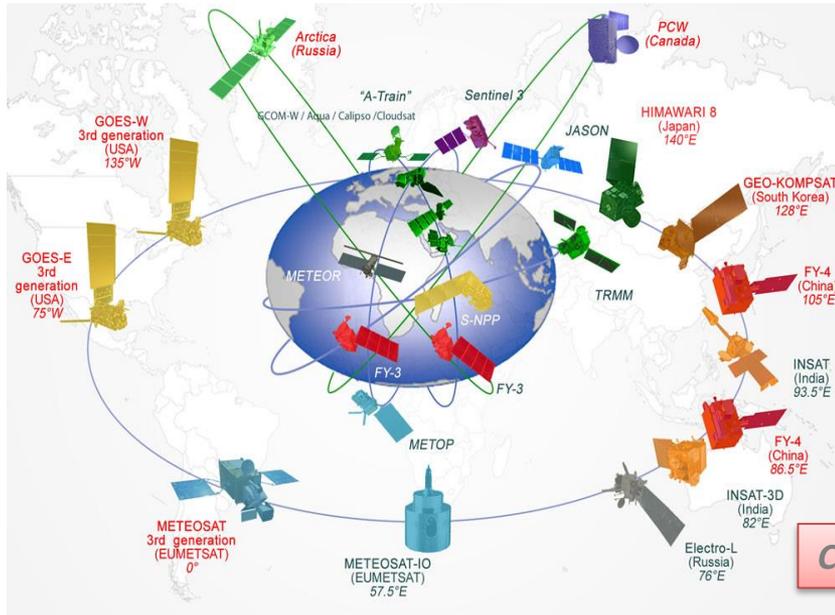
# The Data assimilation cycle



- At ECMWF, twice a day about 30,000,000 observations are used to correct the 150,000,000 variables that define the model's virtual atmosphere.
- This is at ECMWF done by a 4-dimensional adjustment in space and time based on the available observations (4D-Var); this operation takes as much computer power as the 10-day forecast.

# The observations

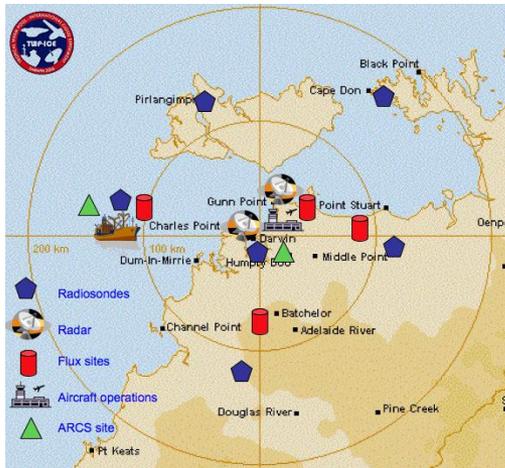
# WMO Integrated Global Observing System



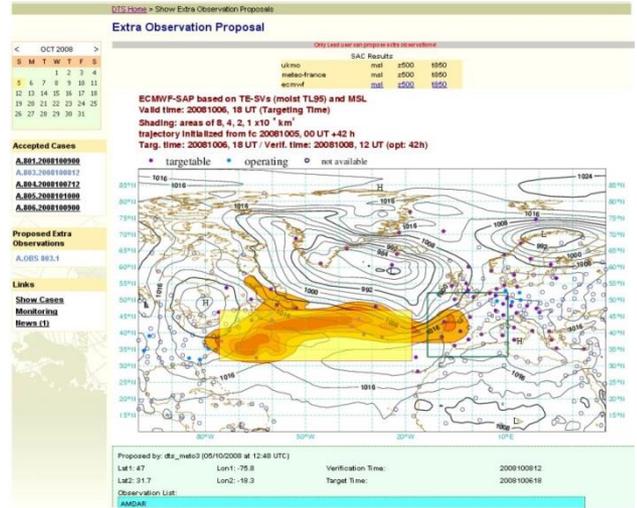
Courtesy: WMO

## ECMWF Preview – DTS

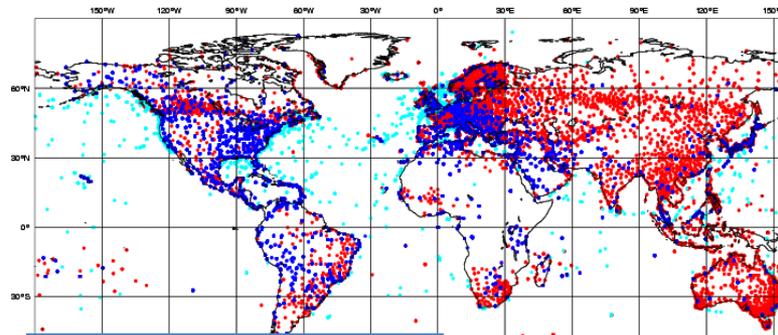
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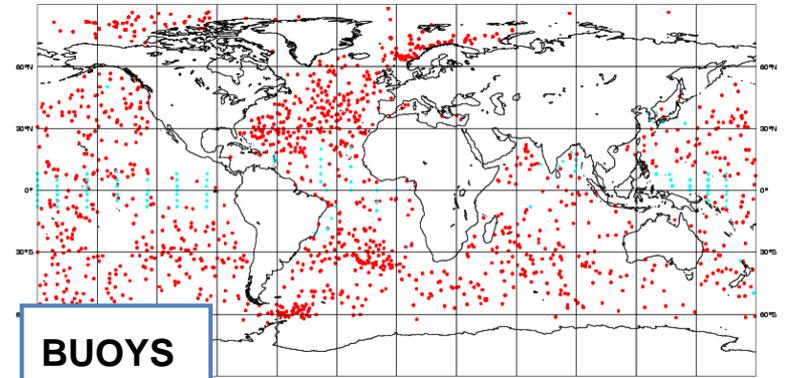
Plus additional obs from field campaign experiments, Data targeting studies, etc.



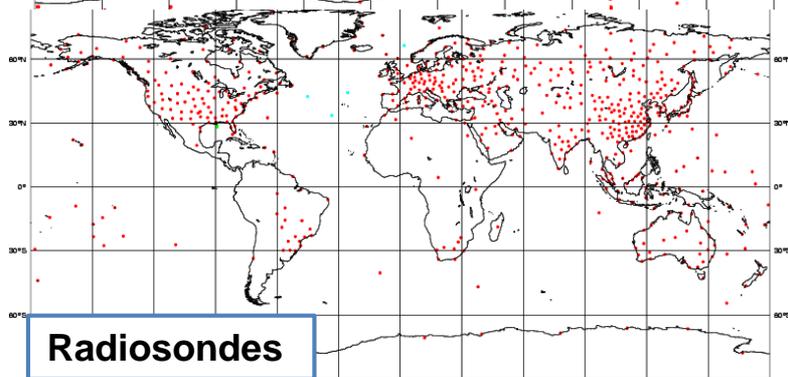
# Distribution of in situ observations



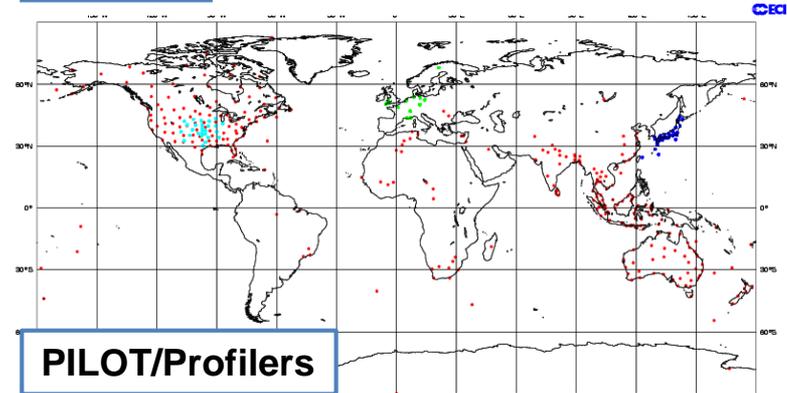
**SYNOP/METAR/SHIP**



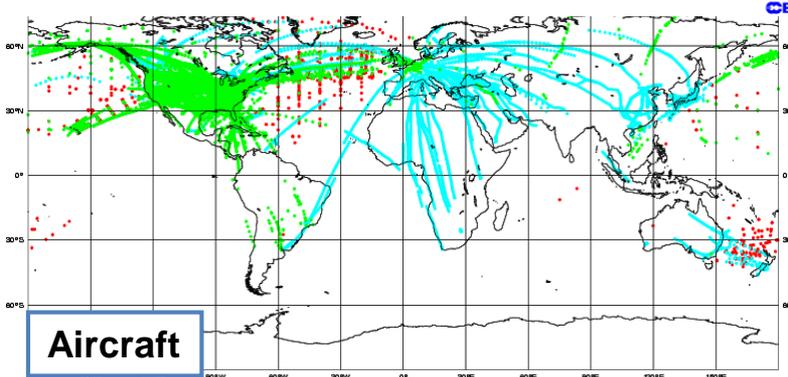
**BUOYS**



**Radiosondes**



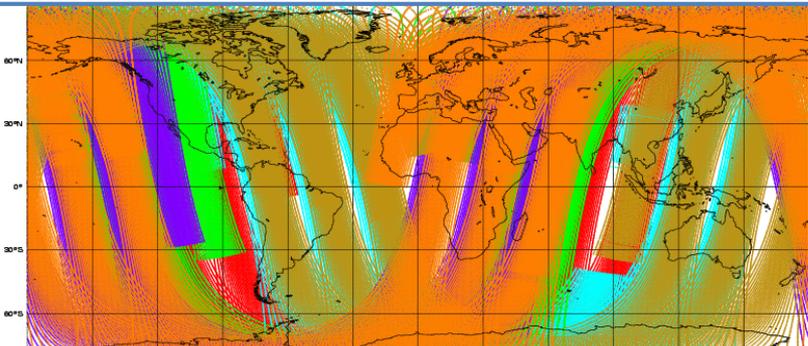
**PILOT/Profilers**



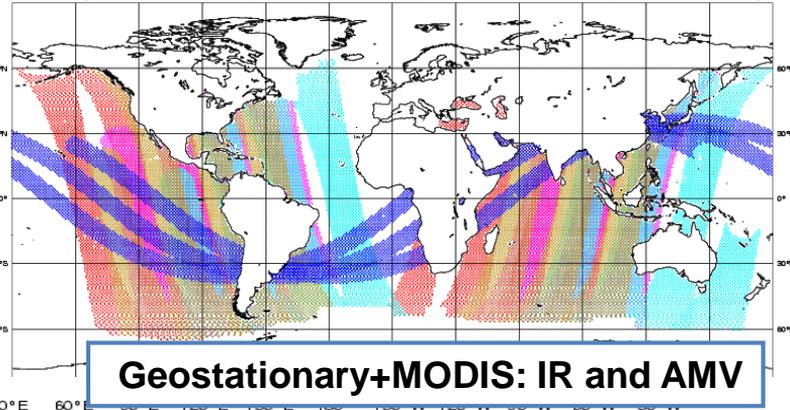
**Aircraft**

# Satellite data sources used by ECMWF's analysis

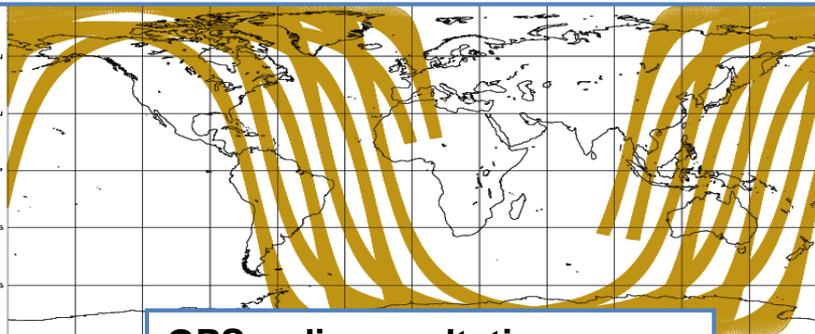
**Sounders: NOAA AMSU-A/B, HIRS, AIRS, IASI, MHS**



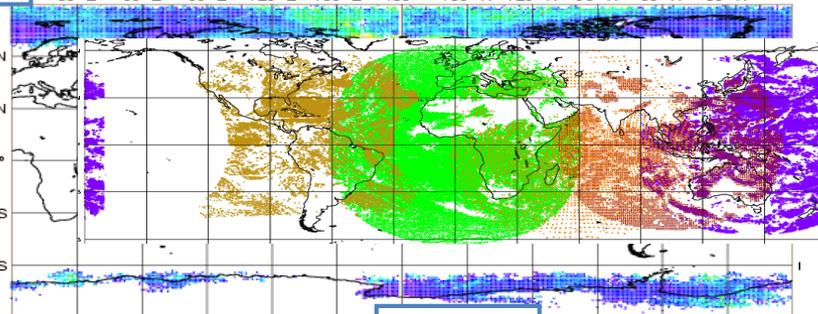
**Imagers: SSMI, SSMIS, AMSR-E,**



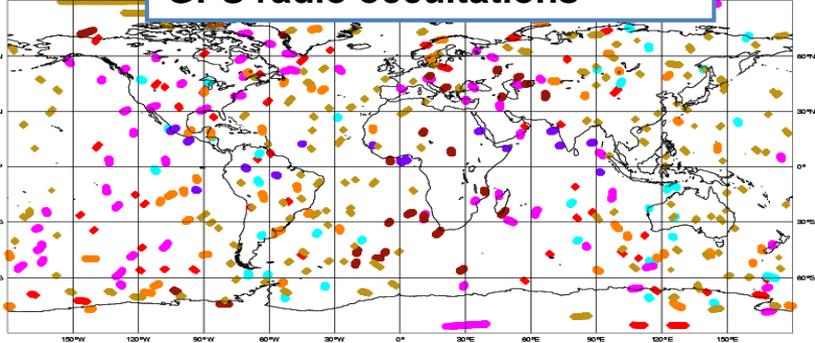
**Scatterometer ocean low-level winds: ASCAT**



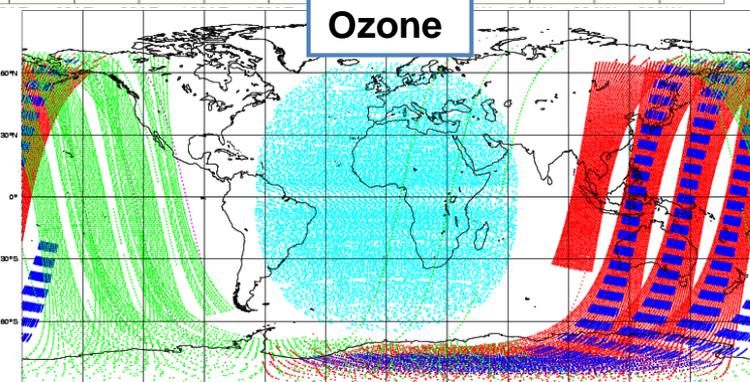
**Geostationary+MODIS: IR and AMV**



**GPS radio occultations**



**Ozone**



# Observation errors

- Even current global observing system is not able to observe the atmosphere completely (data voids): from a mathematical standpoint data assimilation is an under-determined problem
- Most satellite observations (e.g. radiances) are only indirectly related to the quantities of interest (i.e., grid point values of T,u,v,q,O3,...)
- Many satellite observations have coarse vertical resolution
- Most observations measure quantities not located at grid points

In order to compare observations ( $\mathbf{y}$ ) and model ( $\mathbf{x}$ ) we need to perform spatial and temporal interpolations of the model fields and (for satellite observations) transform model fields into radiances: we call this set of operations the **observation operator** ( $\mathcal{H}$ ):

$$\mathbf{y} = \mathcal{H}(\mathbf{x})$$

# Observation errors

- Observations are affected by different types of errors
- Denoting  $\mathbf{y}^*$  as the true observations of the model state ( $\mathbf{y}^* = \mathcal{H}(\mathbf{x}^*)$ ):

$$\mathbf{y} - \mathbf{y}^* = \varepsilon_o = \varepsilon_G + \varepsilon_M + \varepsilon_R + \varepsilon_H$$

$\varepsilon_G =$  **Gross errors** (incorrect coding of observation, duplicates, incorrect location, wrong cloud clearing, etc.).

$\varepsilon_M =$  **Measurement errors** (instrument noise)

$\varepsilon_R =$  **Representativity errors** (e.g., in situ observations compared to grid point model value)

$\varepsilon_H =$  **Observation operator** (Forward model) errors (e.g., errors in the radiative transfer model, interpolation errors, etc.)

# Observation errors

$$\mathbf{y} - \mathbf{y}^* = \varepsilon_o = \varepsilon_G + \varepsilon_M + \varepsilon_R + \varepsilon_H$$

- $\varepsilon_G$  (gross errors) are dealt with by **Observation Quality Control** techniques (more on this later this week); Some of these checks are applied before ingesting the observations (Climatological checks, consistency checks, first guess checks), others are part of the analysis algorithm itself (buddy checks, variational quality control)

- Observations are assumed to be un-biased:

$$\langle \varepsilon_o \rangle = 0$$

- Biases are dealt with specific **Bias Correction** techniques (more on this later this week), which can be part of the analysis algorithm itself (e.g., Variational Bias Correction)

- The covariance matrix of the observation errors is denoted as **R**:

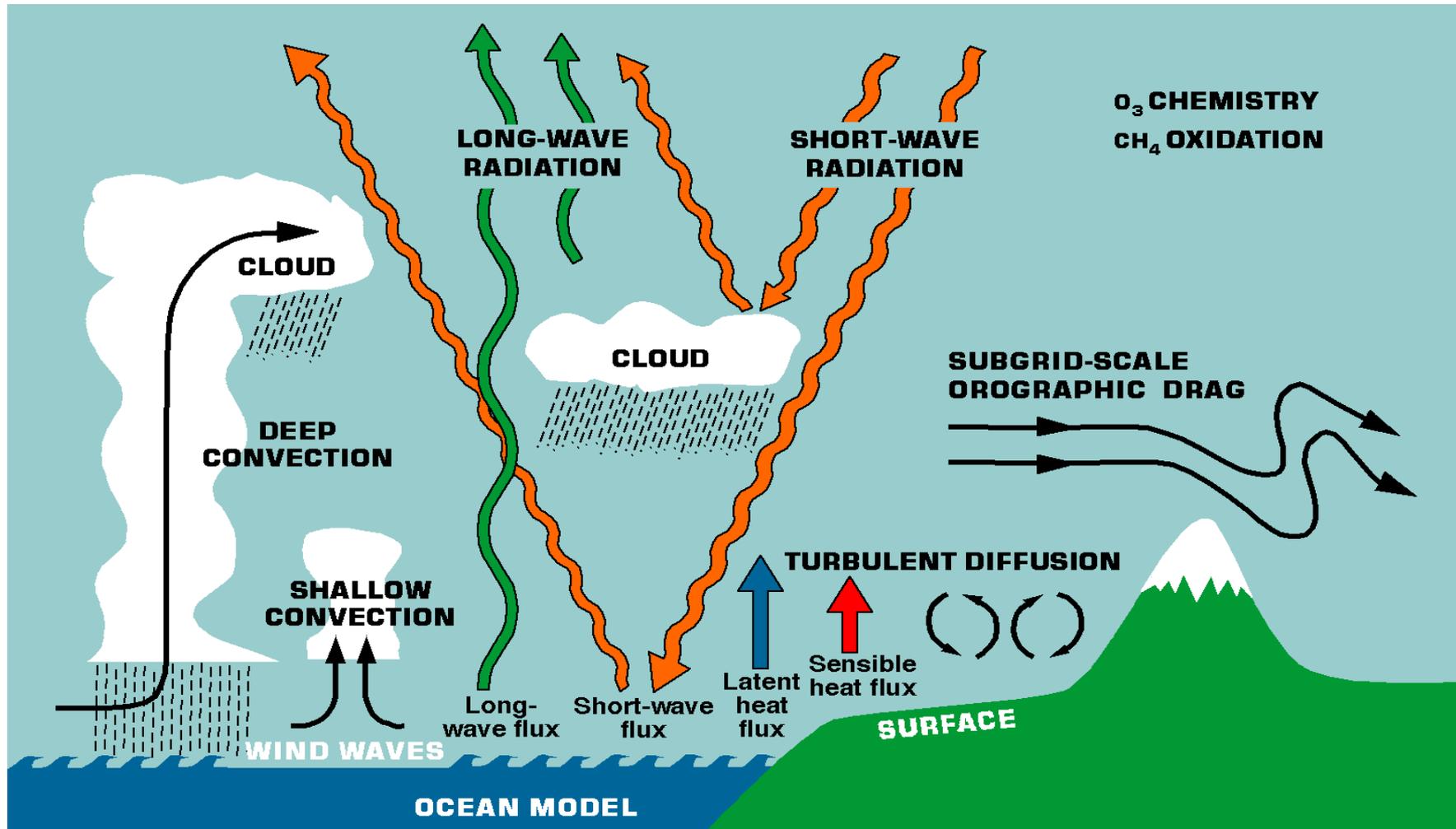
$$\langle \varepsilon_o \varepsilon_o^T \rangle = \mathbf{R}$$

# The forecast model

# The forecast model is a very important part of the data assimilation system

- The short range forecast carries information from past observations into the current analysis (this is called “**the background**”): we need a good model to do this job
- A good model starting from accurate previous analysis will produce an accurate background ➡ the analysis will make only **small corrections** to the background
- In fact when the analysis makes **large corrections** to the background state is usually a sign that something interesting is happening... (e.g., rapid development not present in the forecast; suspect observations)
- In modern data assimilation methods (4D-Var, EnKF, PF) the analysed state is constructed so as to **respect the physical and dynamical balances of the model** ➡ the model is an integral part of the analysis algorithm

# The forecast model is a very important part of the data assimilation system



Most important physical processes in the ECMWF model

# Model errors

- Despite their increasing complexity and sophistication models are far from perfect!
- Sources of model error include: missing physical processes, errors in parametrizations of physical processes, discretisation errors (from continuous PDEs to discrete formulation), etc.,
- We define **model error** as (\* denotes true state,  $i$  is a time index):

$$\mathbf{x}_i^* = \mathcal{M}(\mathbf{x}_{i-1}^*) + \boldsymbol{\eta}_i$$

- Model error can in general have non zero mean:

$$\langle \boldsymbol{\eta}_i \rangle \neq 0$$

- The covariance matrix of the model errors is denoted as  $\mathbf{Q}$ :

$$\langle \boldsymbol{\eta}_i \boldsymbol{\eta}_i^T \rangle = \mathbf{Q}_i$$

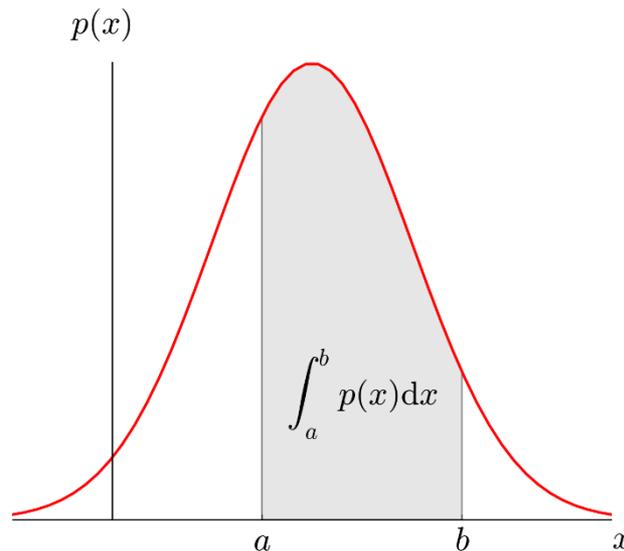
- The treatment of **model error in DA** will be discussed in more detail in a lecture later this week

# Blending observations and model information: the Bayes perspective

# The Bayes perspective

- Both observations and models are affected by random errors
- This means that they should be described as **random variables**
- All we can know about random variables are their **probability distribution functions**:

$$\Pr[a \leq X \leq b] = \int_a^b p(x) dx$$



# The Bayes perspective

- The **Bayes law** descends directly by the definition of conditional probabilities:

$$p(A, B) = p(A|B)p(B) = p(B|A)p(A)$$

⇒

$$p(A|B) = \frac{p(B|A)p(A)}{p(B)}$$

Where:

$p(A, B)$  = probability of events  $A$  and  $B$  happening (**joint** prob. distribution)

$p(A|B)$  = probability of event  $A$  given that  $B$  is true (**conditional** prob. distribution)

$p(A), p(B)$  = probability of event  $A$  ( $B$ ) happening (**marginal** prob. distribution)

# The Bayes perspective

- An illustration ([http://en.wikipedia.org/wiki/Base\\_rate\\_fallacy](http://en.wikipedia.org/wiki/Base_rate_fallacy)):

The police have been issued with breathalysers which never fail to detect a drunk person but have a 5% rate of false positives. Prior campaigns have shown that, on average, one in one thousand drivers drives drunk. If the police stop a driver at random, and he/she results positive to the breathalyser, what is the probability that he/she is actually drunk?

Event A: being a drunk driver. Probability of being a drunk driver, before being tested:  $p(A) = 0.001$

Event B: testing positive to the breathalyser. The probability of testing positive is 1 for the drunken subset of the drivers (0.001) and 0.05 for the sober subset of the drivers (0.999):

$$p(B) = (1 * 0.001) + (0.999 * 0.05) = 0.05095$$

Probability of testing positive to the breathalyser when drunk:  $p(B|A) = 1$

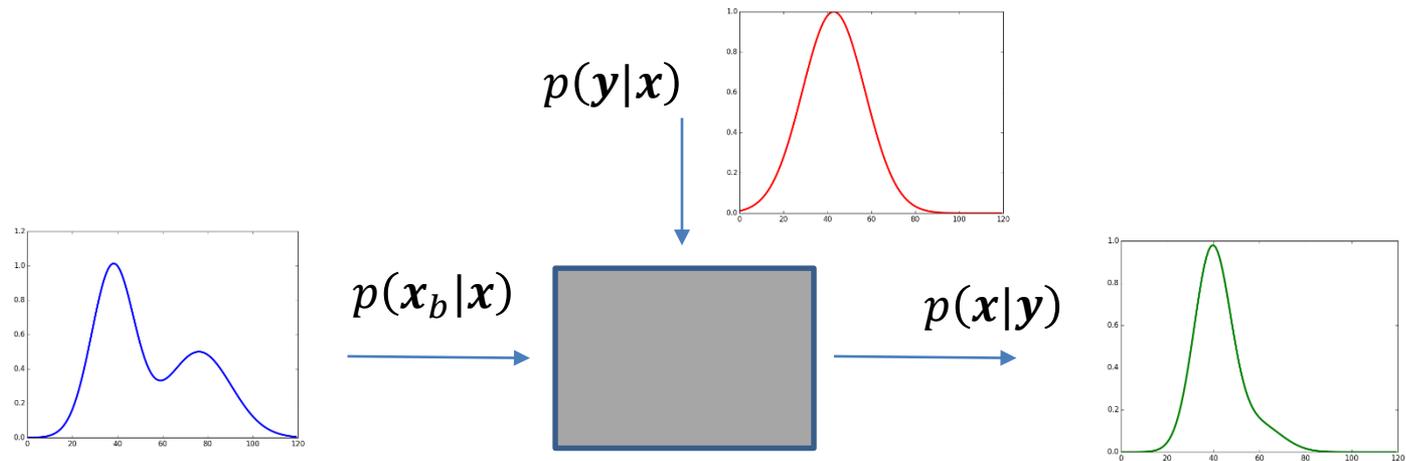
Probability of being drunk after testing positive to the breathalyser,  $p(A|B)$ :

$$p(A|B) = \frac{p(B|A)p(A)}{p(B)} = \frac{1 * 0.001}{0.05095} = 0.0198$$

In words: out of 1000 people stopped by the police, about 51 will result positive, but the probability that anyone of them is actually drunk is less than 2%. (This shows how Bayesian thinking can be useful even beyond data assimilation!)

# The Bayes perspective

- At an abstract level, we can think of the analysis process as updating our prior knowledge about the state, represented by a background forecast and the pdf of its errors, with new observations, represented by their values and the pdf of their errors:



$$p(x|y) = \frac{p(y|x)p(x)}{p(y)} = \frac{p(y|x)p(x_b|x)}{p(y)} \propto p(y|x)p(x_b|x)$$

- $p(x_b|x)$  = **prior pdf** (encapsulate our knowledge about the state before new observations)
- $p(y|x)$  = **observations likelihood** (pdf of the observations conditioned on the state)
- $p(x|y)$  = **posterior pdf** (updated pdf of the state after the analysis)
- $p(y)$  = **marginal pdf of the observations** (does not depend on  $x$ : normalising constant in Bayes' law)

# Particle Filters

$$p(\mathbf{x}|\mathbf{y}) \propto p(\mathbf{y}|\mathbf{x})p(\mathbf{x}_b|\mathbf{x}) \quad (1)$$

- In principle an analysis update requires being able to compute the product pdf of the random variables  $\mathbf{y}$ ,  $\mathbf{x}_b$ . This is usually not possible to do unless we choose very specific functional forms for the pdfs
- We thus need to make approximations
- One idea is to use Monte Carlo methods to sample and propagate the pdfs in (1):  
**Particle Filters**
- In Particle Filters pdfs are sampled by a collection of “particles” (i.e., model states) with assigned weights:

$$p(\mathbf{x}) \sim \sum_{i=1, N} w_i \delta(\mathbf{x} - \mathbf{x}_i) \quad (2)$$

- The pdf is propagated in time by integrating the different particles with the model:

$$p(\mathbf{x}_b|\mathbf{x}) \sim \sum_{i=1, N} w_i \delta(\mathbf{x} - M(\mathbf{x}_i)) \quad (3)$$

# Particle Filters

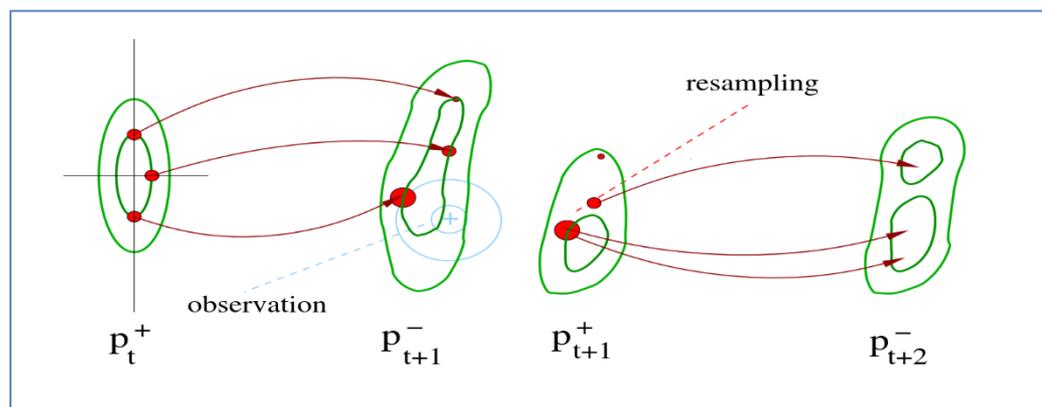
- The pdf is propagated in time by integrating the different particles with the model:

$$p(\mathbf{x}_b|\mathbf{x}) \sim \sum_{i=1, N} w_i \delta(\mathbf{x} - M(\mathbf{x}_i)) \quad (3)$$

- In the analysis update the weights of the particles are updated according to the observations' likelihood:

$$w_i^a \propto w_i p(\mathbf{y}|\mathbf{x}_i)$$

- The ensemble of particles is usually resampled, i.e. high weight particles are duplicated and low-weight particles discarded
- The Particle Filter described here is one of the most basic implementation (Bootstrap Filter, Gordon et al., 1993)



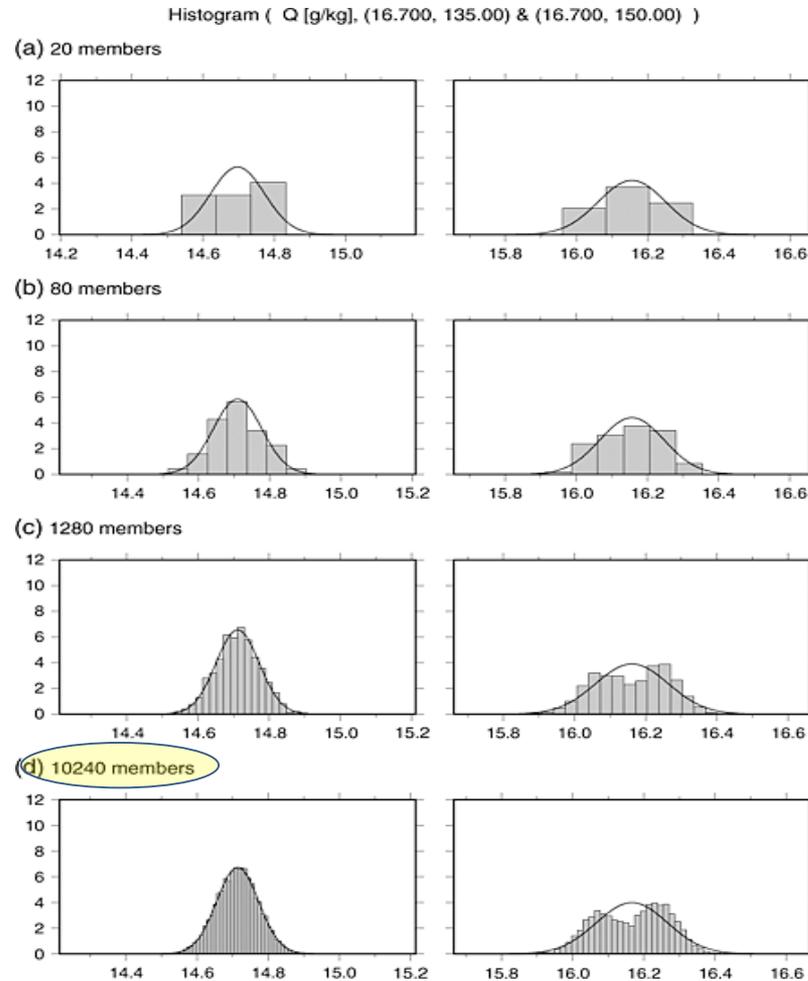
From M. Bocquet

# Particle Filters

- Particle Filters work well for very small state space sizes and observation sizes ( $N \sim 10$  to 100)
- For larger state space and/or observation sizes the required number of particles increases **exponentially** (Snyder et al., 2015)
- A large body of current research is devoted to reduce the computational demands of particle filters for high dimensional systems
- One of the main themes of PF research is how to prevent the particles from diverging from the true state and becoming too unlikely
- The main idea is to also use observations (and not only the model) to “guide” the particles’ evolution from  $t=t_{n-1}$  to  $t=t_n$ ; many variants possible (Ades and van Leeuwen, 2014)

# Particle Filters

- Regardless of the assimilation algorithm the number of particles (ensemble members) needed to resolve non-Gaussian pdfs is very high:



Histograms of a 6 h ensemble forecast for specific humidity ( $\text{g kg}^{-1}$ ) for a intermediate AGCM. Miyoshi et al., 2014

# The Gaussian approximation

- Not making assumptions on the shape of the prior and the likelihood pdf makes the Bayesian problem difficult (i.e., analytically and computationally intractable)
- Usual choice is to assume a **Gaussian distribution** for the both the observations' likelihood and the prior pdf
- Why Gaussian?
  1. Mathematically tractable problem;
  2. Full distribution characteristics defined by only its first two moments (mean and covariance);
  3. Supported by the Central Limit Theorem;
  4. Least committed distribution for given first and second moments (i.e., we are making the least amount of hypotheses on the shape of the pdf)

# The Gaussian approximation

- Usual choice is to assume a **Gaussian distribution** for the both the observations' likelihood and the prior pdf

$$p(\mathbf{y}|\mathbf{x}) = \frac{1}{(2\pi)^{N/2}|\mathbf{R}|^{1/2}} \exp\left(-\frac{1}{2}(\mathbf{y} - \mathbf{H}(\mathbf{x}))^T (\mathbf{R})^{-1}(\mathbf{x}_b - \mathbf{H}(\mathbf{x}))\right)$$

$$p(\mathbf{x}_b|\mathbf{x}) = \frac{1}{(2\pi)^{N/2}|\mathbf{P}_B|^{1/2}} \exp\left(-\frac{1}{2}(\mathbf{x}_b - \mathbf{x})^T (\mathbf{P}_B)^{-1}(\mathbf{x}_b - \mathbf{x})\right)$$

$$p(\mathbf{x}|\mathbf{y}) \propto p(\mathbf{y}|\mathbf{x})p(\mathbf{x}_b|\mathbf{x}) \propto \exp\left(-\frac{1}{2}(\mathbf{y} - \mathbf{H}(\mathbf{x}))^T (\mathbf{R})^{-1}(\mathbf{y} - \mathbf{H}(\mathbf{x})) - \frac{1}{2}(\mathbf{x}_b - \mathbf{x})^T (\mathbf{P}_B)^{-1}(\mathbf{x}_b - \mathbf{x})\right)$$

- where  $\langle \varepsilon_o \varepsilon_o^T \rangle = \mathbf{R}$  and  $\langle \varepsilon_b \varepsilon_b^T \rangle = \mathbf{P}_B$  are the covariances of the errors of the observations and of the prior (background forecast)
- Under this assumption the posterior (analysis) distribution  $p(\mathbf{x}|\mathbf{y})$  can also be expressed as a Gaussian distribution

# Kalman Filter methods

- Once we know (at least in principle!) the form of the posterior distribution  $p(\mathbf{x}|\mathbf{y})$  we have a choice:
  - 1) Either we can solve for the mean and the covariance of the posterior distribution:

$$\mathbf{x}_a = \int \mathbf{x} p(\mathbf{x}|\mathbf{y}) d\mathbf{x}$$

$$\mathbf{P}_a = \int (\mathbf{x} - \mathbf{x}_a)(\mathbf{x} - \mathbf{x}_a)^T p(\mathbf{x}|\mathbf{y}) d\mathbf{x}$$

Methods based on this approach include Optimum Interpolation (O.I.), Kalman Filter, Ensemble Kalman Filter (EnKF). These will all be discussed this week. The analysis found through this approach is referred to as the **minimum variance** solution or the best linear unbiased estimate (**BLUE**).

Note: Kalman Filter based methods can be derived without making any assumptions about the Gaussianity of the errors. However only if all error distributions are Gaussian will the KF provide the exact posterior distribution (i.e. the Bayes posterior pdf).

# Variational methods

- 2) Alternatively we might choose to estimate the mode of the posterior distribution  $p(\mathbf{x}|\mathbf{y})$ , i.e. find the analysis  $\mathbf{x}_a$  as the state that corresponds to the maximum of the posterior distribution ( $\Rightarrow$  the most probable state):

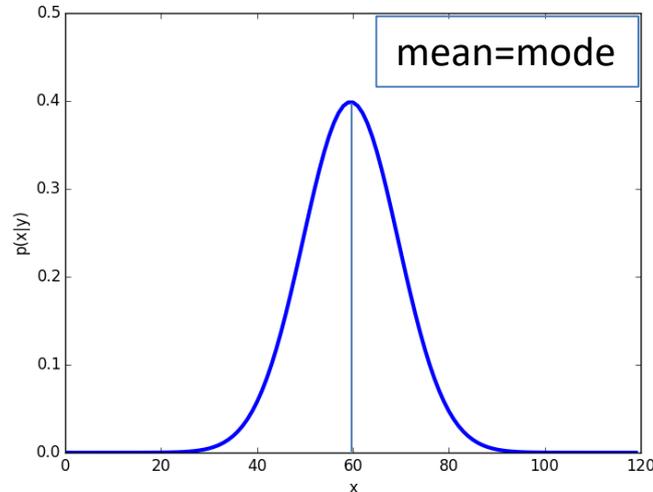
$$\mathbf{x}_a = \arg \max_x (p(\mathbf{x}|\mathbf{y}))$$

This way of attacking the problem leads to the variational approach (3D-Var, 4D-Var). They will be covered extensively in this week's lectures. The solution found in this way is called the **maximum a-posteriori probability** state (**MAP**).

In the variational framework the linear and Gaussian assumptions can be relaxed, i.e. the full nonlinear analysis problem can be decomposed into a series of linear Gaussian problems (incremental 4D-Var, to be discussed later this week). However there is no guarantee of convergence!

# Kalman Filter vs Variational methods

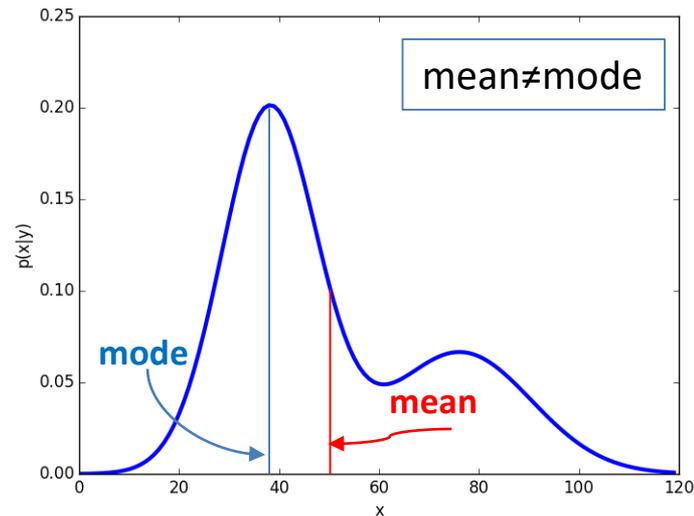
- For a Gaussian pdf the mean and the mode coincide:



- Thus if all the relevant statistics are Gaussian the minimum variance and maximum a-posteriori solutions coincide

# Kalman Filter vs Variational methods

- For non-Gaussian pdfs the mean and the mode of the distribution generally differ:



- In non-Gaussian assimilation problems the minimum variance and maximum likelihood solutions will differ

# Hybrid DA methods

- Both Variational and Kalman Filter based analysis methods require estimates of the background state and its error covariances ( $p(\mathbf{x}_b|\mathbf{x}) \sim \mathcal{N}(\mathbf{x}_b, \mathbf{P}_B)$ )
- The background state is provided by an integration of the prognostic model started from the previous analysis:

$$\mathbf{x}_b^t = \mathcal{M}(\mathbf{x}_a^{t-1})$$

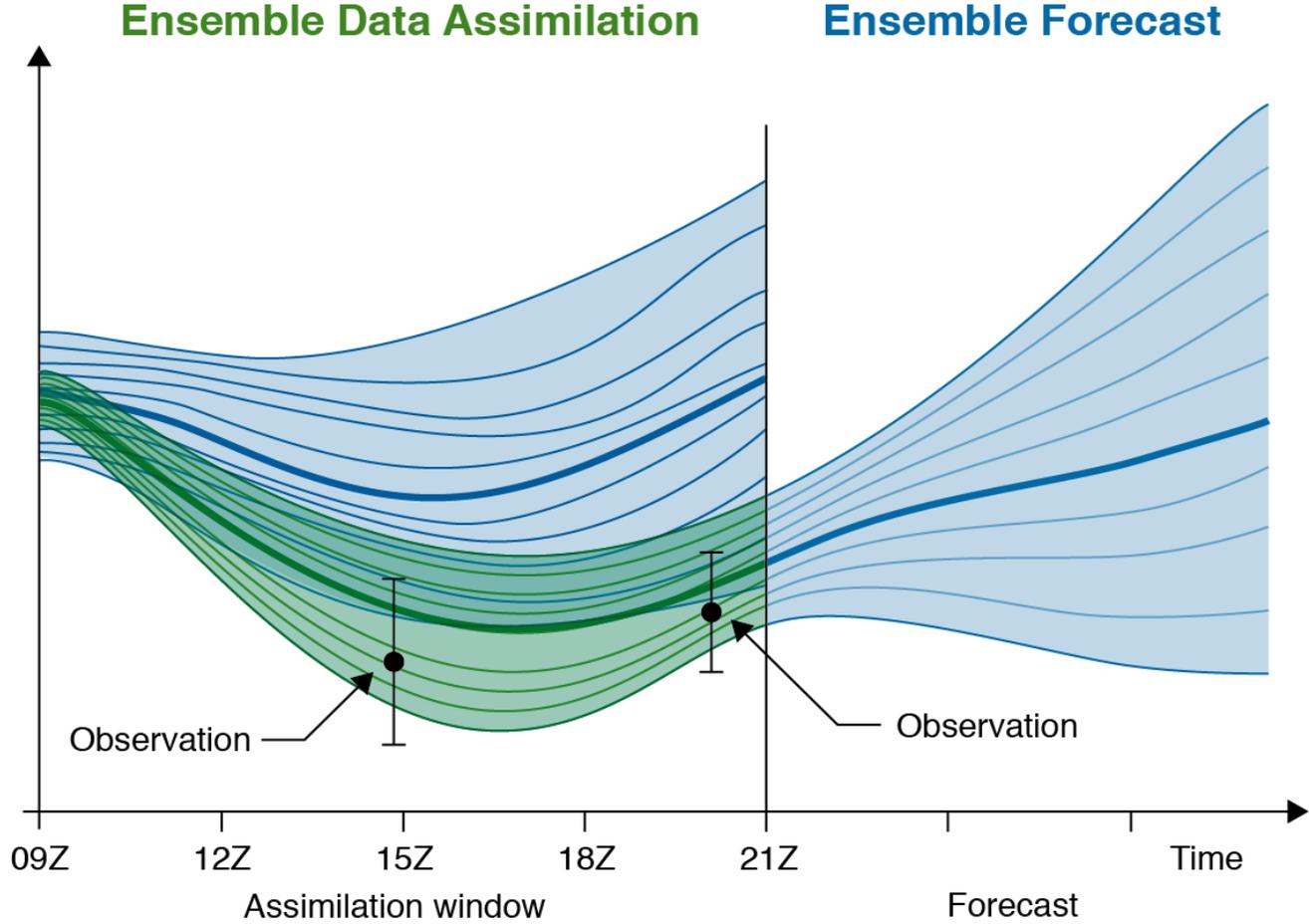
- The background (and analysis) error statistics are usually computed with **Monte Carlo methods**: an ensemble of states is used to estimate the errors statistics
- Each ensemble member is advanced using a perturbed version of the model:

$$\mathbf{x}_{b,i}^t = \mathcal{M}(\mathbf{x}_{a,i}^{t-1}) + \boldsymbol{\eta}_i \quad i = 1, 2, \dots, N_{ens}$$

- Each ensemble member is usually updated using perturbed observations (though there are methods that can avoid this):

$$\mathbf{y}_i = \mathbf{y} + \boldsymbol{\varepsilon}_{o,i} \quad i = 1, 2, \dots, N_{ens}$$

# Hybrid DA methods



(E. Källén)

# Hybrid DA methods

- Data assimilation systems that have an ensemble data assimilation component used for the estimation of analysis and background errors are called **hybrid data assimilation systems** (Note: this definition is not universal!)
- We will discuss the various options for the ensemble data assimilation component in two dedicated lectures
- All major global NWP Centres run some form of hybrid data assimilation for atmospheric DA: a variational analysis cycle to estimate the mean/mode of the analysis pdf coupled with an ensemble data assimilation system to estimate the second moments of the distributions.
- The ensemble DA component not only serves the purpose of estimating the background errors used in the analysis update, but it also provides the initial state estimates for ensemble forecasting

# Summary

- Data assimilation in NWP aims to optimally blend information from observations and model to produce an accurate and physically consistent estimate of the initial state of the atmosphere and of the other components of the Earth System
- Both observations and models are affected by systematic and random errors: these need to be evaluated and taken into account in order to produce a statistically optimal analysis
- The Bayesian approach provides a unified theoretical framework for data assimilation
- Particle Filters provide a Monte Carlo implementation of the Bayes' Law in data assimilation. Asymptotically correct for  $N_{ens} \rightarrow \infty$ , but not yet applicable to high dimensional systems
- A Gaussian assumption on the error statistics is usually made to make the problem tractable in realistic geophysical DA
- Kalman Filter type methods and Variational methods can both be derived from Bayes' Law under these assumptions: they lead to the same solution for linear, Gaussian problems
- Hybrid data assimilation methods currently used in global NWP combine a variational analysis system with an ensemble data assimilation component for error estimation

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