

Use of SMOS data in a coupled land-atmospheric model

sensitivity to different model and observation error scenarios

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P. de Rosnay, C. Albergel, G. Balsamo, L. Isaksen, M. Drusch and many other

SMOS & ECMWF

➤ **Mission objective:** Provide global measurements of two key variables in the water cycle: soil moisture and ocean salinity.

L-band mission (innovative 2D interferometric radiometer); transparent to clouds, large penetration depth, less sensitive to vegetation canopy and soil roughness.

➤ **Objectives at ECMWF:**

- Global **monitoring** of T_B at the satellite antenna reference frame, in NRT
- Assimilation of SMOS T_B over continental surfaces & investigate the meteorological impact of SMOS data assimilation
- Introducing new observations is an efficient way to improve the forecast/analysis

Monitoring SMOS TB

➤ **Routinely production of statistics with SMOS T_B , model equivalents and background departures, in NRT**

- Global scale
- Land and oceans separately,
- Several incidence angles [10, 20, 30, 40, 50, 60],
- Two polarisations states [XX, YY],
- Independently per continents and hemispheres,

➤ **Statistical products,**

- Time-averaged geographical mean-fields (last 6 weeks of data),
- Hovmöller zonal mean fields (last 3 months),
- Time series of area averages (last 3 months),
- Angular distribution of bias: background departures as function of incidence angle (last 5 weeks).

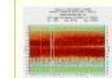
➤ **Support to CAL/VAL sites** → time series produced for 17 sites

564 images are produced and updated daily → important contribution to the SMOS quality control

Time-averaged geographical mean fields



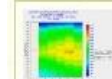
Hovmoeller zonal mean fields



Time series of area averages



Scatter plots



Time series of area averages (over land)



Time series of area averages (over Sea)



Time series of targeted sites statistics (over Land)



SMOS & ECMWF

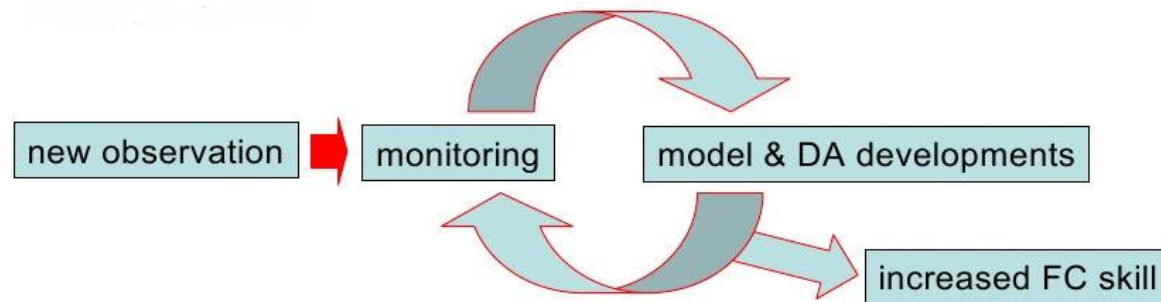
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Soil moisture analysis at ECMWF

Simplified Extended Kalman Filter:

For each grid point, analysed state vector \mathbf{x}_a :

$$\mathbf{x}_a = \mathbf{x}_b + \mathbf{K} (\mathbf{y} - \mathcal{H}[\mathbf{x}_b])$$

\mathbf{x}_b : background state vector,

\mathbf{y} : observation vector

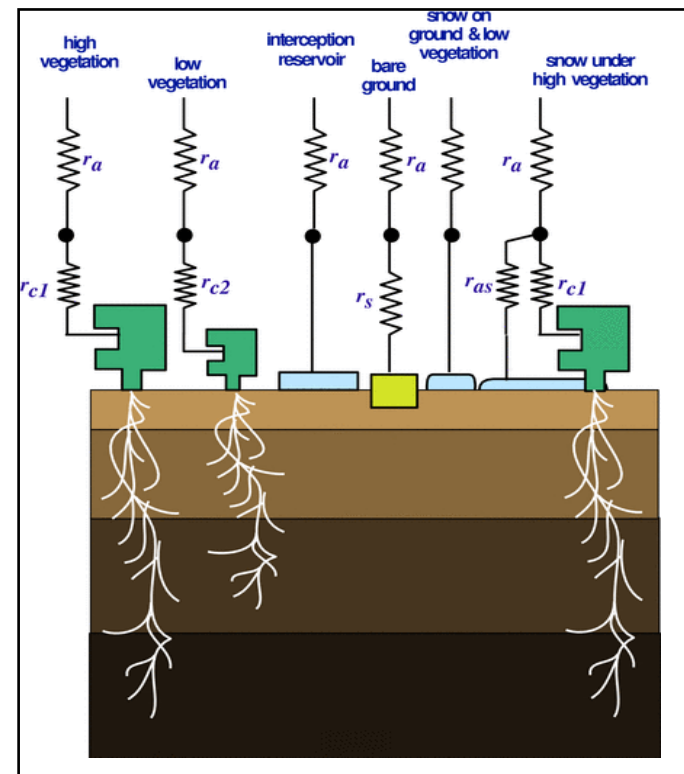
\mathcal{H} : non linear observation operator

\mathbf{K} : Kalman gain matrix

$$\mathbf{K} = [\mathbf{B}^{-1} + \mathbf{H}^T \mathbf{R}^{-1} \mathbf{H}]^{-1} \mathbf{H}^T \mathbf{R}^{-1}$$

Observations:

- Operations: screen level variables (SLV): T^{2m} , RH^{2m}
- Research:
 - ASCAT soil water index (METOP-A, METOP-B),
 - SMOS Brightness temperatures

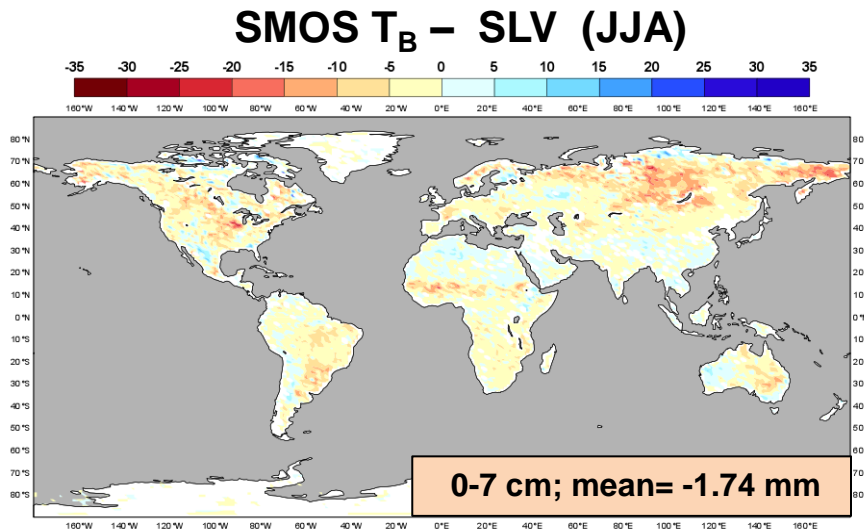


LSM : H-TESSEL

0-7cm, 7-28cm, 28-100cm,
100-289cm

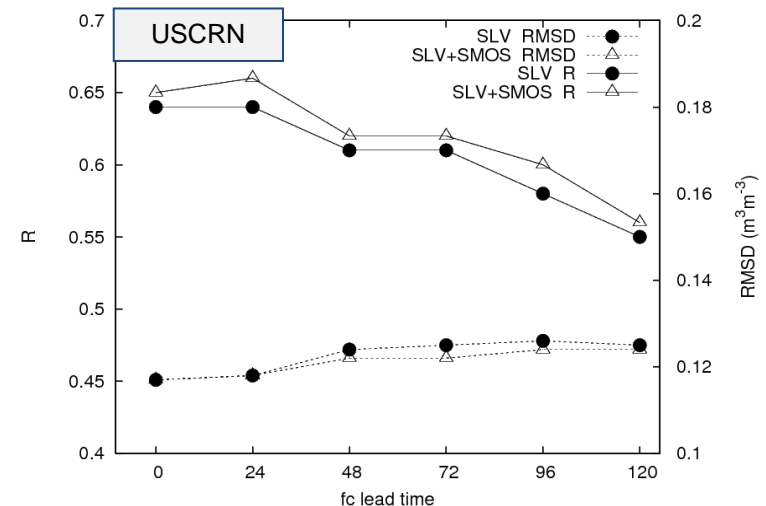
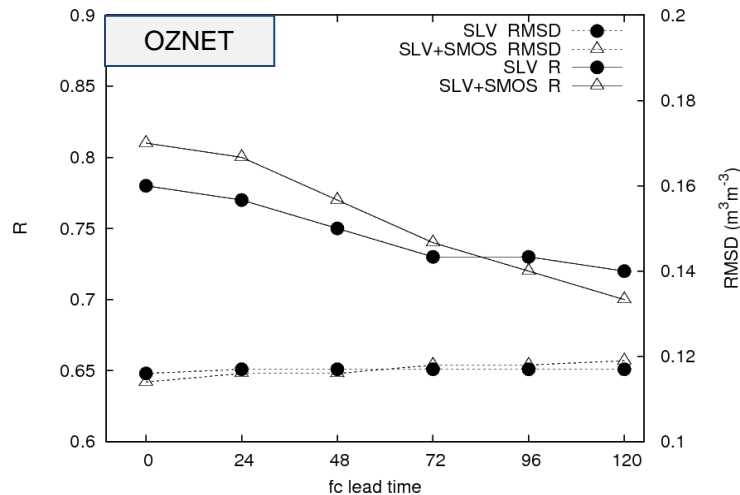
(Balsamo et al., JHM, 2009)

SMOS DA impact experiments



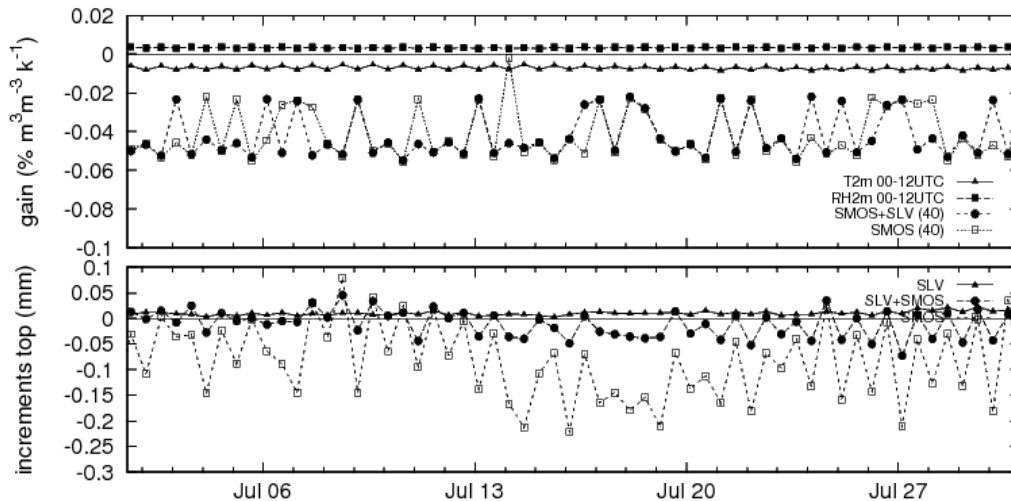
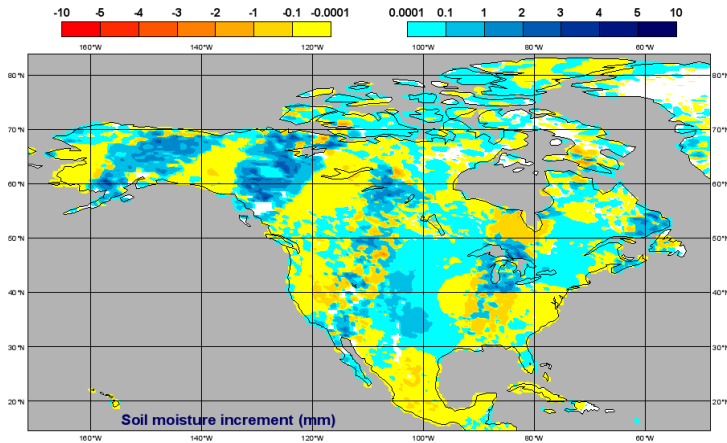
SM analyses were validated against more than 600 in-situ stations in 10 different countries:

- Impact on soil moisture is high!,
- SM dynamic is improved and bias reduced,
- Root-zone is better characterised,
- Skill in the forecast of soil moisture is kept at least up to 72 h.



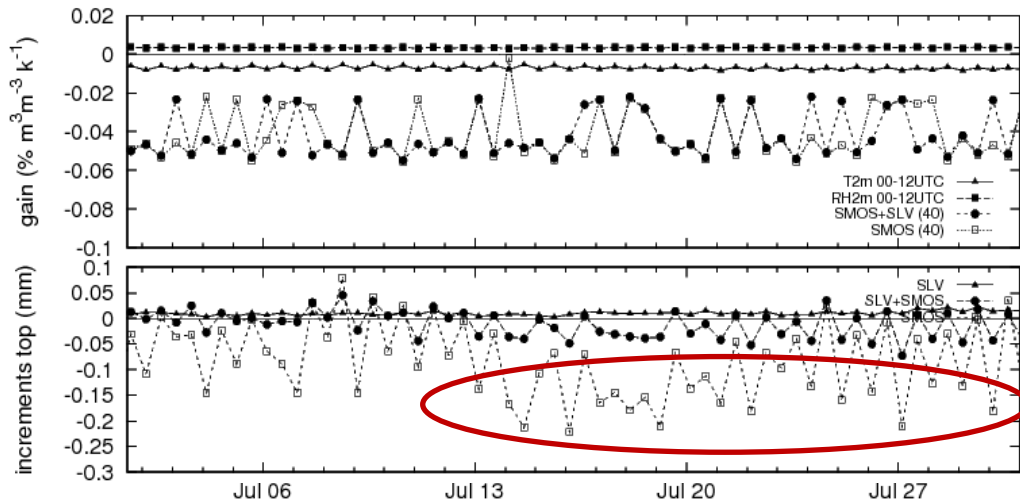
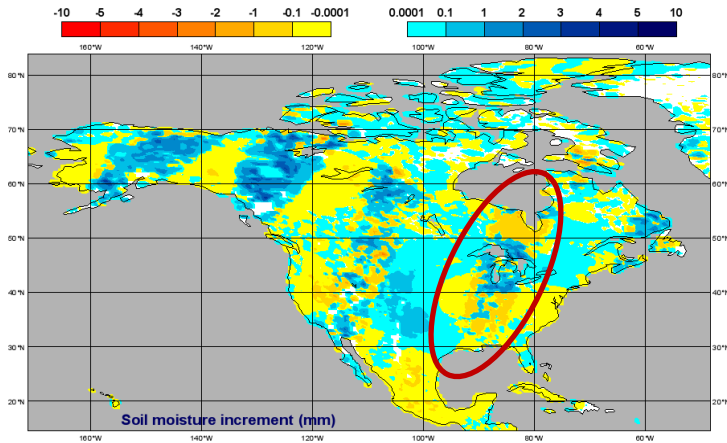
SMOS DA impact experiments

Assimil T_{2m}, RH_{2m} & SMOS T_B



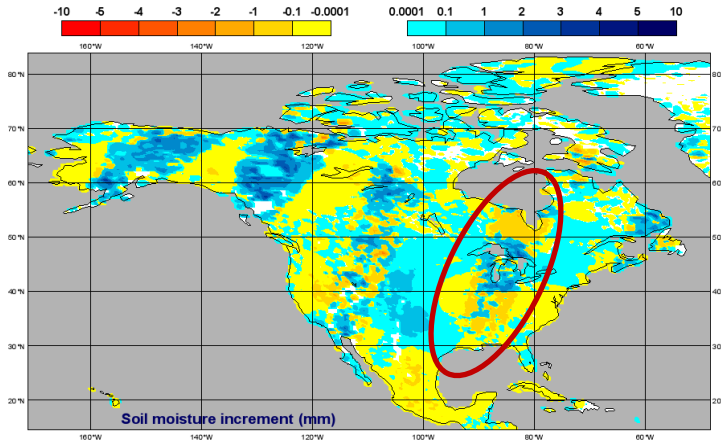
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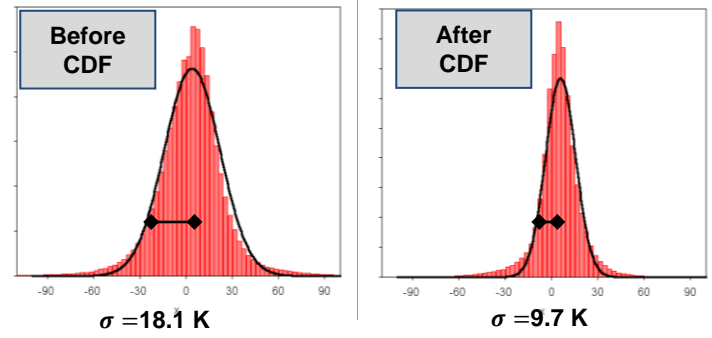
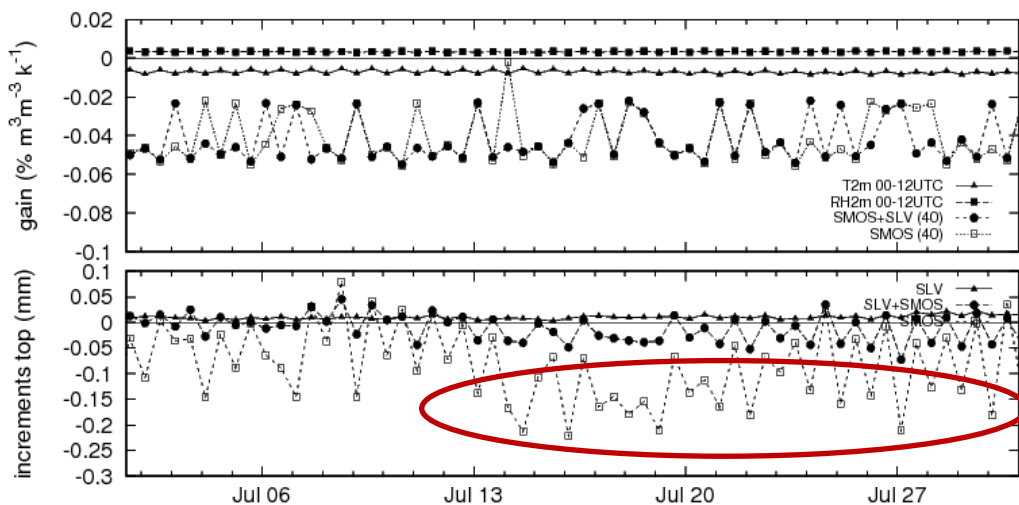


SMOS DA impact experiments

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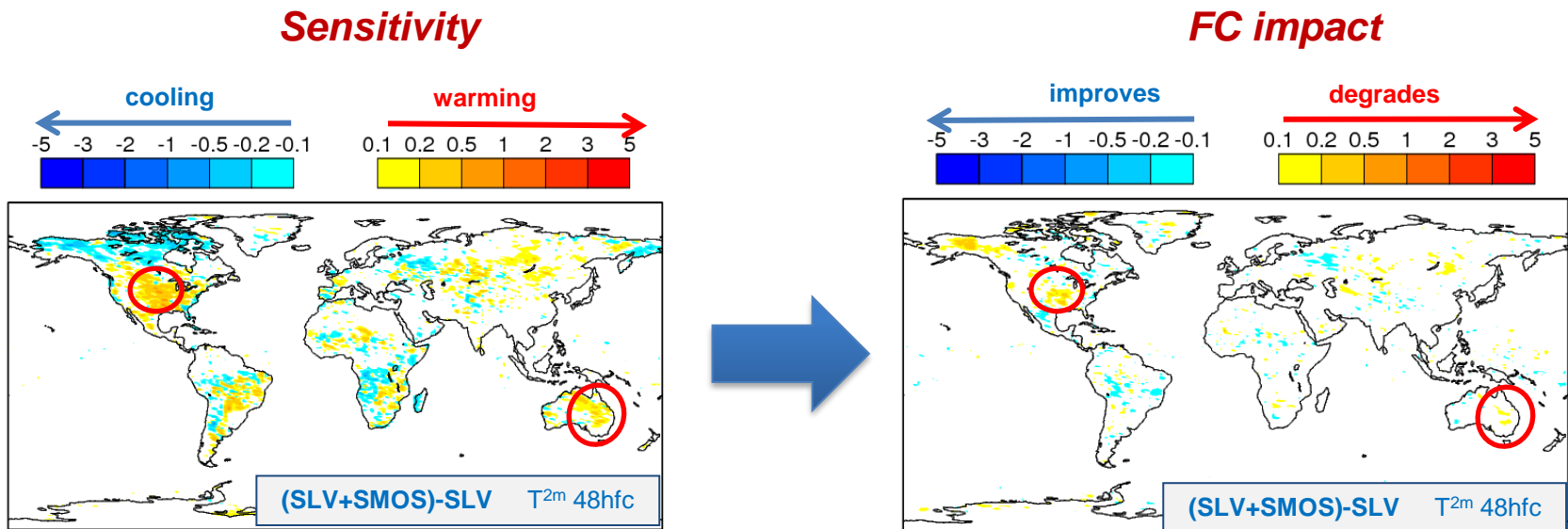
Large sensitivity → large Gain
 +
 Residual large bias
 =
 Large increments



$$K = [B^{-1} + H^T R^{-1} H]^{-1} H^T R^{-1}$$

Gain (and increments) very sensitive to error covariance matrices

Impact in the forecast skill



→ SMOS increments produce warmer and drier atmosphere in center US, Sahel, South of Africa and Australia → hot-spots for NWP impact,

→ Small impact in the skill of the forecast by assimilating SLV+SMOS.

Sensitivity experiments

Investigate the effect of various types of assimilated observations, assimilation approach and observation (**R**) and background error (**B**) specification in the soil moisture analysis.

- USA → best place for availability of observations and “cheaper” experiments,
- Period: 15 Sept- 14 Oct 2012 → recharge period, good variability of soil moisture,
- Full coupled system,
- 3 angles (30, 40, 50), 2 polarisations (XX, YY), AF-FOV, RFI flag,
- Physics of cy40r1,
- Reduced observing system for the upper-air atmosphere; ATOVS, GBRAD and NEXRAD observations used to limit number of observations, and still reasonable atmospheric constrain
- **R** cov matrix: $\sigma(T^{2m}) = 2 \text{ K}$; $\sigma(RH^{2m}) = 10\%$; $\sigma(\text{SMOS } T^B) = \text{rad_acc K}$
- **B** cov matrix: $\sigma(\text{sm}_{(0-7) \text{ cm}}) = \sigma(\text{sm}_{(7-28) \text{ cm}}) = \sigma(\text{sm}_{(28-100) \text{ cm}}) = 0.01 \text{ m}^3\text{m}^{-3}$
- **Q** cov matrix: $\sigma(\text{sm}) = 0.01 \text{ m}^3\text{m}^{-3}$

Experiment types

- **OL** → free soil moisture run,
- **SLV** → assimilation of only T^{2m} , RH^{2m} (simulate surface operational conditions)
- **SLV+SMOS** → assimilation of T^{2m} , RH^{2m} and SMOS T_B with **B** static
- **SMOS** → assimilation of only SMOS T_B with **B** static
- **SMOS PDI** → pseudo direct-insertion of SMOS T_B . SEKF filters still apply to increments and departures
- **SMOS 2R** → assimilation of only SMOS T_B , doubling the observation error (2R),
- **SMOS B-prop** → assimilation of only SMOS T_B with **B** propagated between two cycles. Background error grows along the assimilation window
- **SMOS B-text** → assimilation of only SMOS T_B with **B** propagated between two cycles. Background error grows along the assimilation window. For a medium texture soil, 10% of WHC is equivalent to doubling background error (0.02 m^3m^{-3}), or 20 mm for the 1st meter of soil.
- **SMOS 3DB** → an 3D structure background error is assumed. The model top layer is more affected by short term variability and more sensitive to precipitation errors → 20% of WHC for top layer (~ 0.04 m^3m^{-3} for medium-type soil), 10% of WHC for 2nd layer and 5% of WHC for 3rd more stable layer.



Type of assimilated observation

Experiment types

- OL → free soil moisture run,
- SLV → assimilation of only T^{2m} , RH^{2m} (simulate surface operational conditions)
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- **SMOS B-prop** → assimilation of only SMOS T_B with **B** propagated between two cycles. Background error grows along the assimilation window. Model error was set to $0.01 \text{ m}^3\text{m}^{-3}$,
- **SMOS B-text** → assimilation of only SMOS T_B ; background error is defined as a proportion of the water holding capacity (WHC). For a medium-type soil (medium WHC), 20 mm for the 1st m, 10 mm for the 2nd m, and 5 mm for the 3rd m (0.02 m^3m^{-3}), or 20 mm for the 1st m, 10 mm for the 2nd m, and 5 mm for the 3rd m
- **SMOS 3DB** → an 3D structure background error is assumed. The model top layer is more affected by short term variability and more sensitive to precipitation errors → 20% of WHC for top layer ($\sim 0.04 \text{ m}^3\text{m}^{-3}$ for medium-type soil), 10% of WHC for 2nd layer and 5% of WHC for 3rd more stable layer.

Weight given to SMOS observations

Experiment types

Different B matrix structures

• OL → free soil moisture run,

• SLV → assimilation of only T^2m , R

• SLV+SMOS → assimilation of T^2m , R

• **SMOS B-fix** → assimilation of only SMOS T_B with **B** static

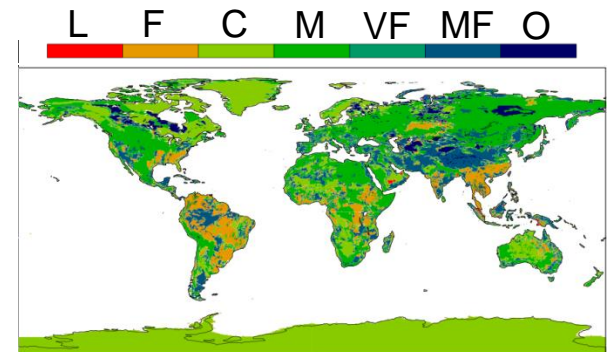
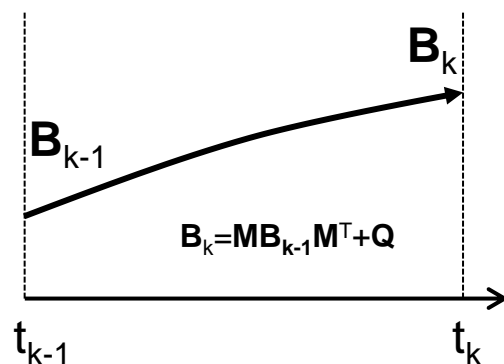
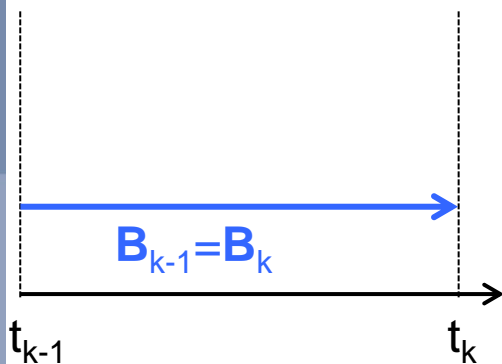
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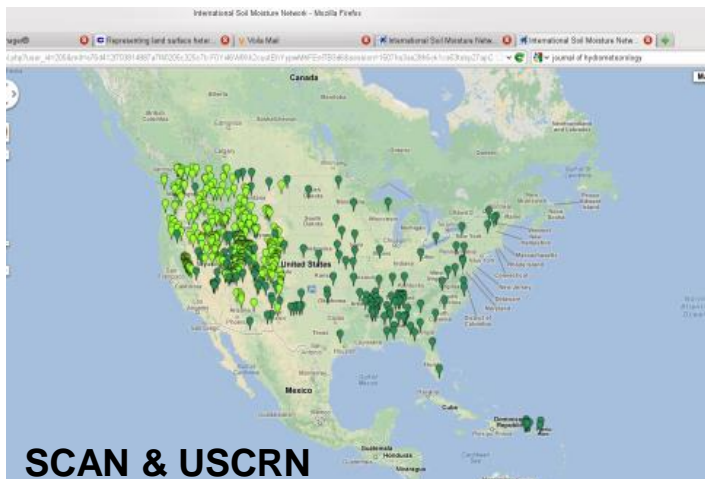
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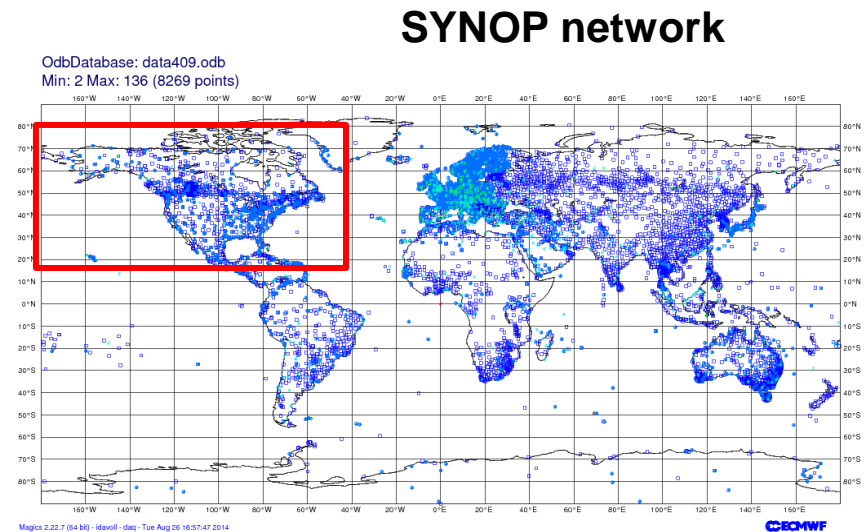


Validation and verification

- Validation against in-situ soil moisture data from two independent networks: SCAN and USCRN
- Comparison against 2 m temp and 2 m dew point temp observations from the SYNOP network
- Atmospheric verification using a North-America mask



<http://ismn.geo.tuwien.ac.at/ismn/>

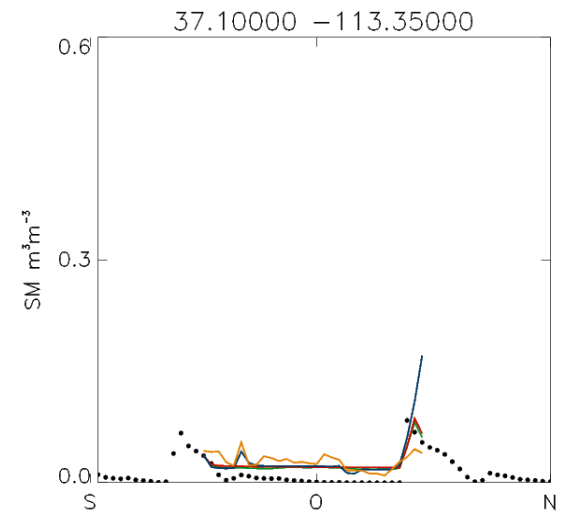
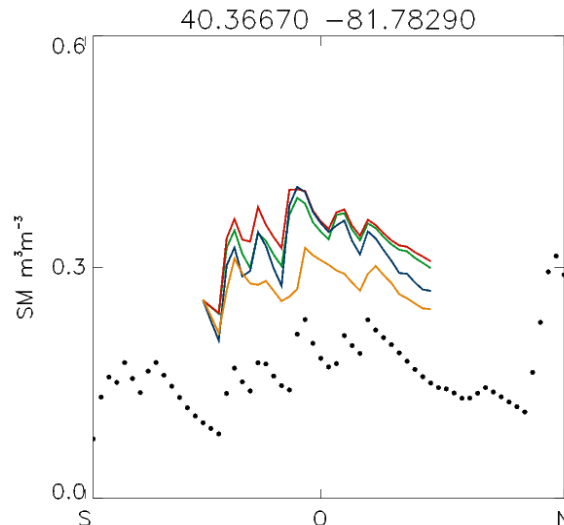
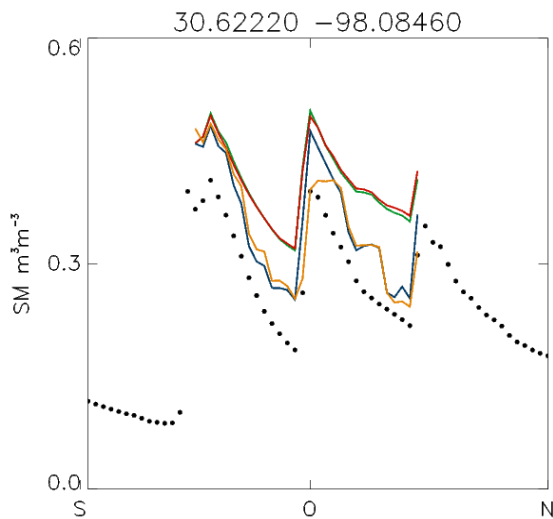


Validation against in-situ data

USCRN	Bias (m^3m^{-3})	RMSD (m^3m^{-3})	R	N	SCAN	Bias (m^3m^{-3})	RMSD (m^3m^{-3})	R	N
OL	-0.115	0.130	0.75	60	OL	-0.062	0.104	0.74	86
SLV	-0.115	0.130	0.75	60	SLV	-0.061	0.104	0.74	86
SMOS+SLV	-0.097	0.121	0.76	60	SMOS+SLV	-0.048	0.101	0.75	86
SMOS	-0.089	0.115	0.67	60	SMOS	-0.035	0.101	0.68	86

Only stations with significant correlation values Confid. Int. 95% (p-value < 0.05)

— SMOS + SLV
— SMOS



Atmospheric scores

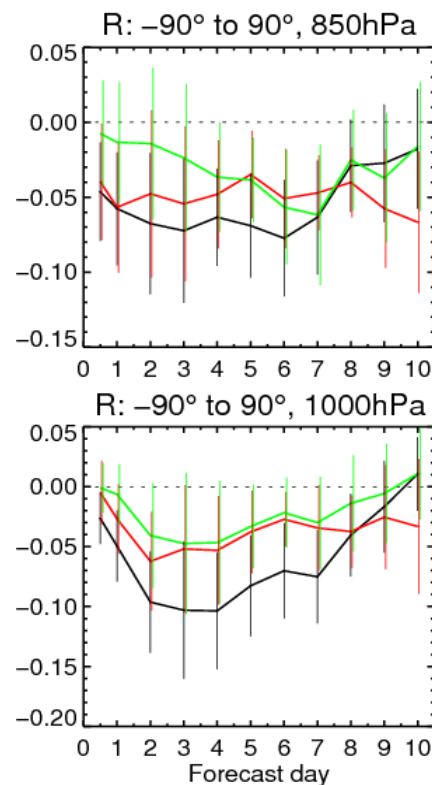
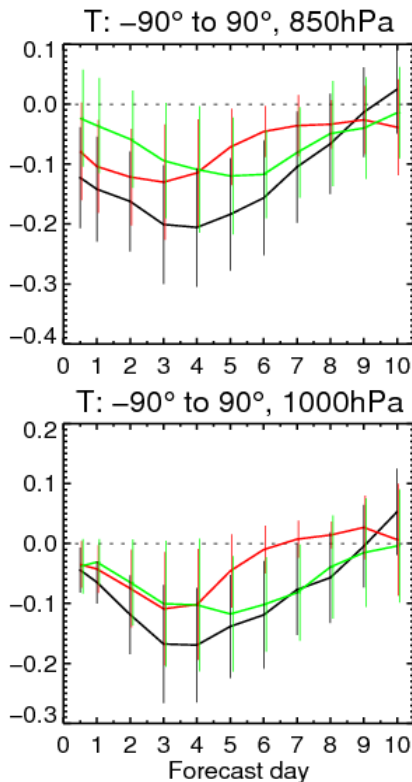
Normalized change in rms of fc error: $drmse = \frac{RMS(e^{SMOS}) - RMS(e^{CTRL})}{RMS(e^{CTRL})}$; $e^{EXPT} = fc^{EXPT} - an^{REF}$
 $e^{CTRL} = fc^{CTRL} - an^{REF}$

_____ SLV - OL
 _____ SMOS - OL
 _____ SLV+SMOS - OL

$drmse > 0 \rightarrow$ expt increases error
 $drmse < 0 \rightarrow$ expt decreases error

Temperature

Humidity



Atmospheric scores

Normalized change in rms of fc error: $drmse = \frac{RMS(e^{SMOS}) - RMS(e^{CTRL})}{RMS(e^{CTRL})}$; $e^{EXPT} = fc^{EXPT} - an^{REF}$
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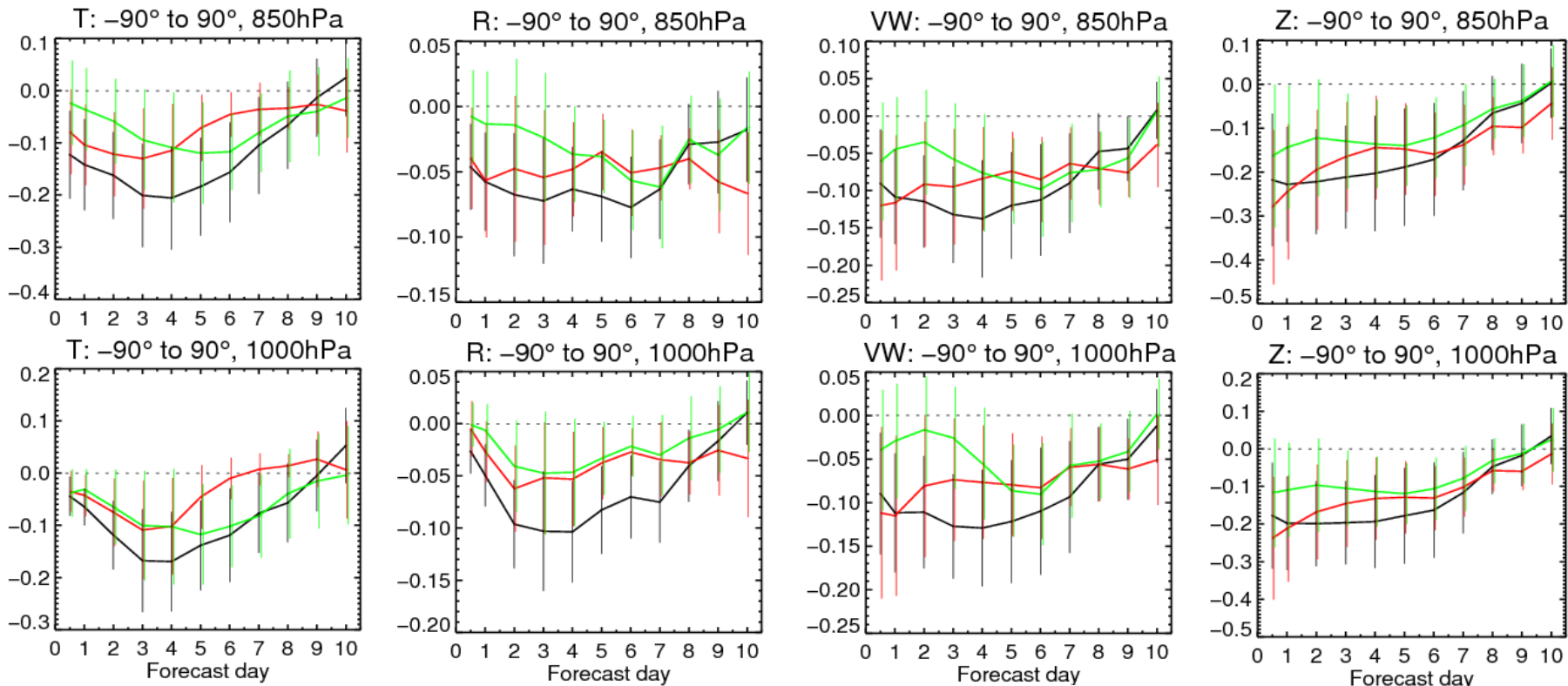
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Humidity

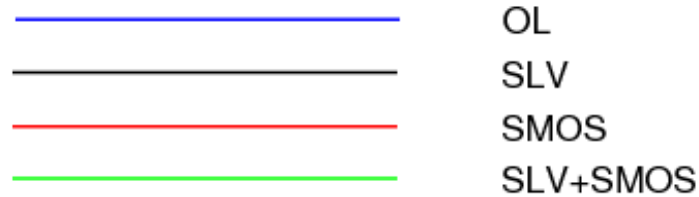
Vector wind

Geopotential

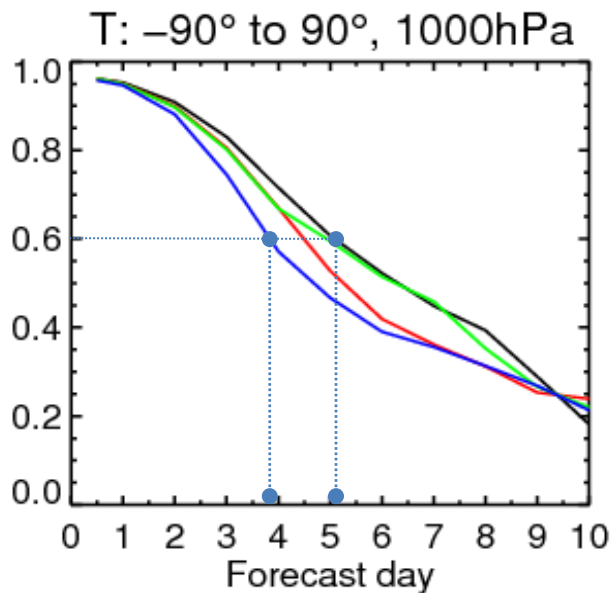


Atmospheric scores

Anomaly correlation

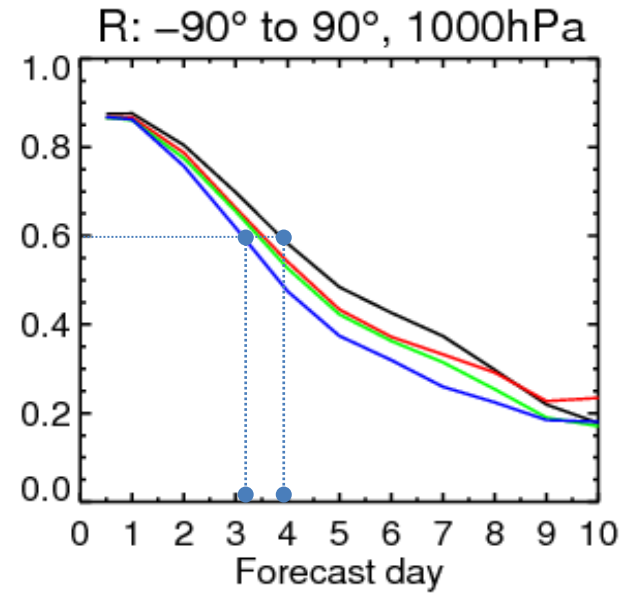


Temperature



> 1 day

Humidity



~ 1 day

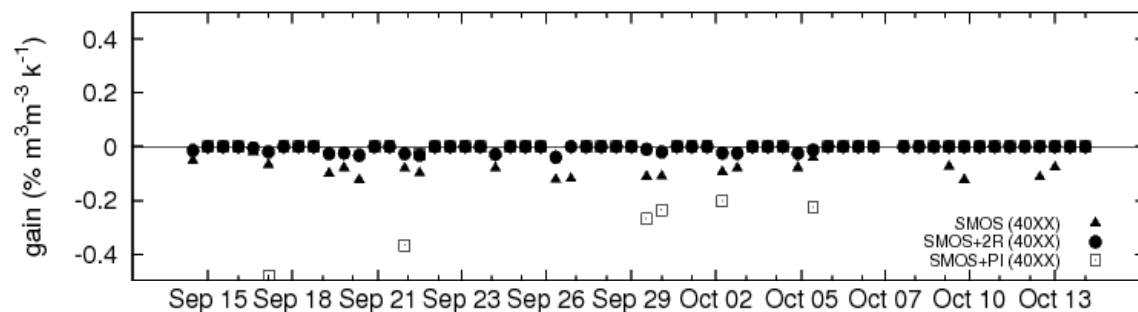
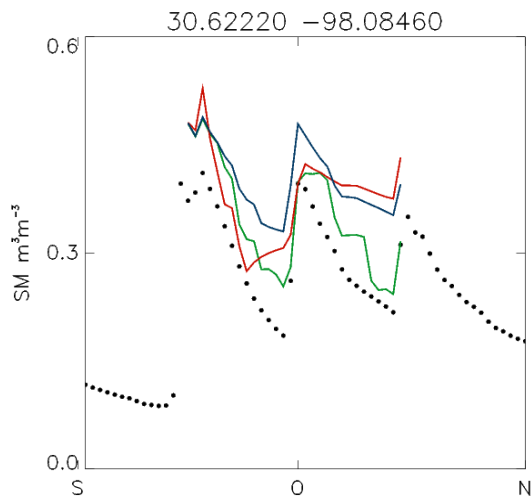
Validation against in-situ data

USCRN	Bias (m^3m^{-3})	RMSD (m^3m^{-3})	R	N
Direct Ins	-0.099	0.116	0.71	58
SMOS + R	-0.086	0.113	0.69	58
SMOS+2R	-0.096	0.117	0.74	58

SCAN	Bias (m^3m^{-3})	RMSD (m^3m^{-3})	R	N
Direct Ins	-0.051	0.106	0.68	83
SMOS + R	-0.032	0.101	0.69	83
SMOS+2R	-0.044	0.104	0.70	83

*Only stations with significant correlation values
Confidence 95% (p-value < 0.05)*

— SMOS R
— SMOS 2R



→ Good impact of SMOS+2R in the root-zone (R)

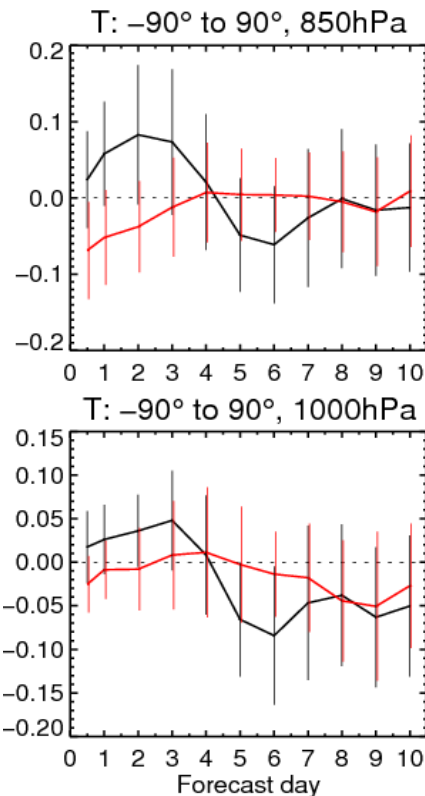
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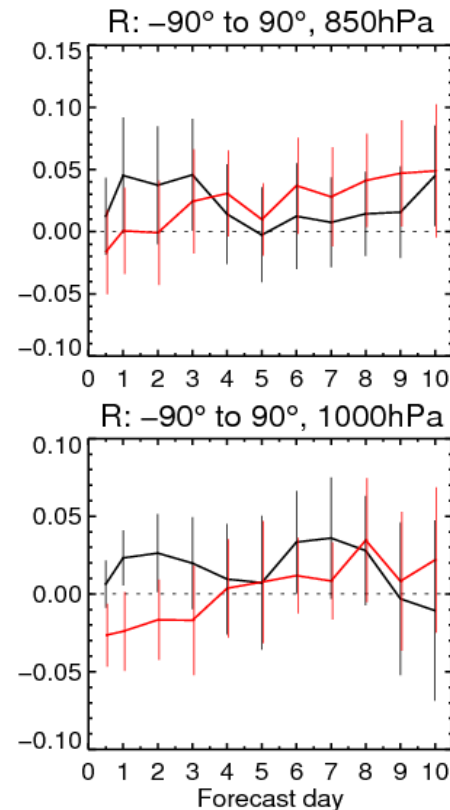
— SMOS+PI – SMOS
 — SMOS+2R – SMOS

drmse > 0 → expt increases error
 drmse < 0 → expt decreases error

Temperature



Humidity



Validation against in-situ data

USCRN	Bias (m ³ m ⁻³)	RMSD (m ³ m ⁻³)	R	N
SMOS B-fix	-0.085	0.109	0.70	64
SMOS B-prop	-0.088	0.111	0.69	64
SMOS Btext	-0.074	0.104	0.67	64
SMOS + 3DB	-0.071	0.102	0.65	64

SCAN	Bias (m ³ m ⁻³)	RMSD (m ³ m ⁻³)	R	N
SMOS B-fix	-0.022	0.095	0.70	77
SMOS B-prop	-0.025	0.095	0.70	77
SMOS Btext	-0.015	0.094	0.66	77
SMOS + 3DB	-0.016	0.094	0.64	77

*Only stations with significant correlation values
Confidence 95% (p-value < 0.05)*

→ **Low impact in the root-zone**

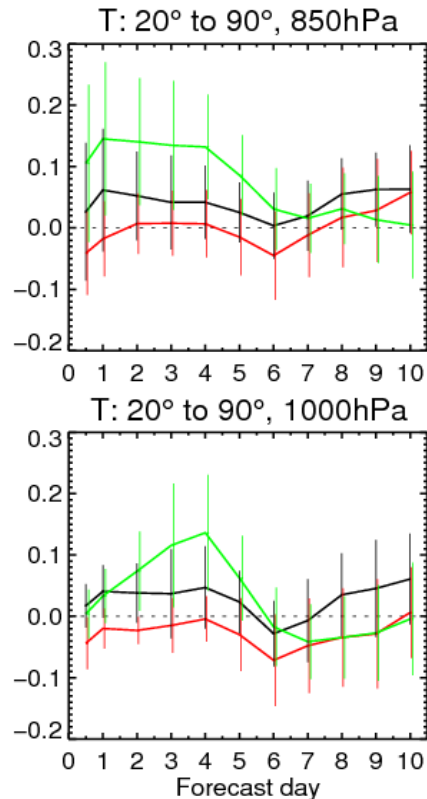
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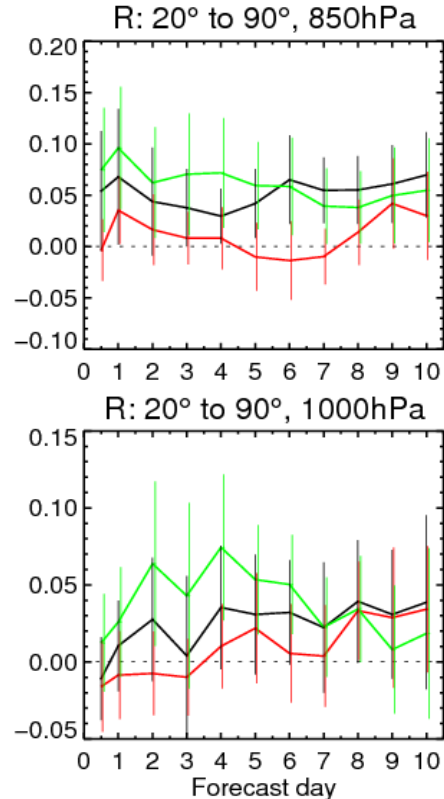
- SMOS B-text - SMOS
- SMOS 3DB - SMOS
- SMOS B-prop - SMOS

$drmse > 0 \rightarrow$ expt increases error
 $drmse < 0 \rightarrow$ expt decreases error

Temperature



Humidity



Conclusions (I)

- SMOS data successfully integrated into a coupled land-atmospheric model,
- First seasonal experiments show that, compared to the operational system, the SMOS signal tend to dry the soil
 - positive results in terms of shallow and root-zone soil moisture,
 - Possible compensation mechanism in the atmosphere,
- Several diagnostics show that several components of the assimilation system should be adjusted to optimize the use of SMOS information in the land DA system,
- Sensitivity experiments:
 - G-I: Type of observation:
 - Constraining soil moisture through observations is important,
 - Soil moisture analyses benefit of assimilating SMOS data,
 - But main improvement of atmospheric variables produced by SLV (improvement up to 20% and 1-week)
 - Compensation mechanisms in coupled system

Conclusions (II)

➤ G-II: Weight of SMOS observation:

- Given total confidence to observations produces spurious increments,
- Doubling SMOS observation error does not reduce RMSD for top layer, but improves the correlation and slightly the atmospheric scores → fair increase of the observation error

➤ G-III: B-matrix error structures:

- Introducing soil texture information in the background error is beneficial for soil moisture,
- Atmospheric scores are neutral

→ Doubling SMOS observation error and introducing soil texture information in the background error, in combination with SLV, could improve land and atmospheric scores,

→ We are still learning!

Thanks for your attention !

contact: joaquin.munoz@ecmwf.int

Further information:

SMOS online monitoring in NRT:

<http://www.ecmwf.int/products/forecasts/d/charts/monitoring/satellite/smos/>

ECMWF SMOS website:

http://www.ecmwf.int/research/ESA_projects/SMOS/index.html

ECMWF CMEM website:

http://www.ecmwf.int/research/data_assimilation/land_surface/cmем/cmем_index.html

Questions

- How is possible that sm is barely affected by assimilating SLV observations, and however it has a great impact in the atmosphere compared to the OL?
 - XXX (locally changes and degradations ??)
- Why when SMOS data is assimilated, the bias against in-situ is greatly reduced, but why not the RMSE?
 - XXX
- Should we use anomaly correlation as alternative metric?
 - Yes, I think we shouldn't to complement the metric observed here and to quantify the skill of SMOS data to predict short-scale variability.
- Why have you used the operational analysis as reference and not the own analysis?
 - Because the oper offers the best possible analysis as reference. In our case upper-air observing system is reduced as I wanted faster experiments, and the quality of the analysis are likely to be not as good as those of the oper.
- Why B-prop doesn't add any improvement?
 - Because the B matrix is not cycled but propagated during 12 h and reinitialized at the next cycle. The benefits of propagating the B-matrix are over a time scale of 3 and 9 hours, as the analysis are at 00 and 06 UTC. For these scales and the error given to Q, errors do not grow much and little impact is observed,